Rate- and State-based Simulations of Induced Seismicity and Coupling to Reservoir Processes

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Simulation Goals

- Perform simulations of induced events
- Develop a method to determine safe injection practices
- Inform regulatory agencies
- Simulations must produce realistic results
 - Power law FMD
 - Migration of events away from well with time
 - Small events that lead to larger events
 - M_{max} scales with injected volume (?)
 - Low stress drops
 - Rupture speeds on the order of traditional VEQ

What information is necessary to simulate induced earthquakes?

- 1. Fault geometry and rate-state constitutive parameters
- 2. Reservoir characterization
- 3. External stressing history
 - Analytical solution for pore-fluid diffusion (Wang, 2002)
 - NUFT (Nitao, 1998; Hao et al., 2012)
- 4. Tectonic driving stress (perhaps neglect this in relatively aseismic regions)
- 5. Pre-existing shear stress conditions:
 - In situ stress measurements regional average (from global stress maps)
 - Projection of the regional stress tensor (from global stress maps)
 - Randomly generated heterogeneous field (some fractal distribution)
- 6. Earthquake simulation method
 - RSQSim
- 7. Well located seismicity catalog with low magnitude of completeness

Earthquake Time-dependency: Rate- and State-dependent Friction

$$\tau = \mu(\sigma - p)$$
$$\mu = \mu_0 + a \ln\left(\frac{V}{V^*}\right) + b \ln\left(\frac{\theta V^*}{D_c}\right)$$

 $\tau = (\sigma - p) \left| \mu_0 + \frac{\sigma}{V} \ln\left(\frac{V}{V^*}\right) \right| \mathbf{b} \ln\left(\frac{\theta V^*}{D_0}\right) \right|$

(1) Modified Coulomb Criterion

(2) Rate- and State-dependent friction

(3) Constitutive Law

Rate-term State-term

- μ_0 = Nominal coefficient of friction
- V*: Reference slip rate
- V: Earthquake slip rate
- θ : State variable
- *D_c*: Characteristic slip distance

a and b: Constitutive parameters describing the material

RSQSim: Governing equations

• Constitutive relation:

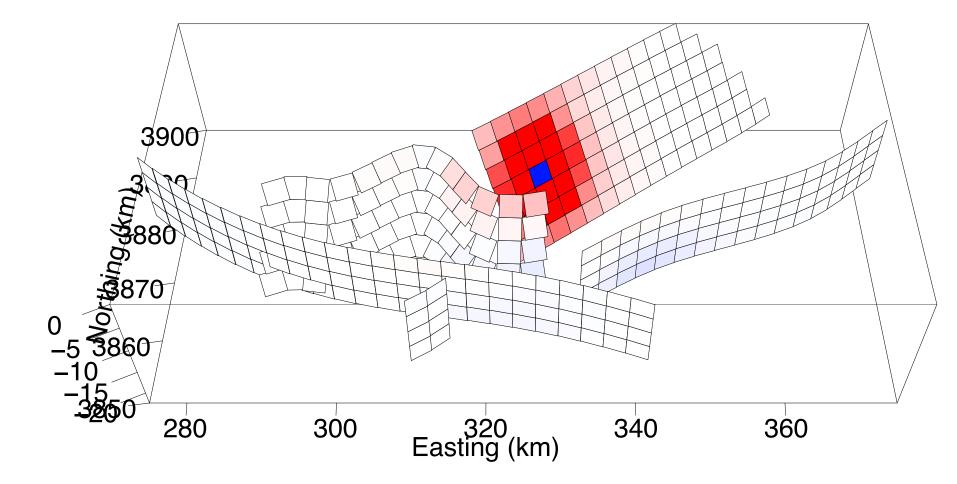
$$\tau_i^{\text{fric}} = \sigma_i \left[\mu_0 + a \ln\left(\frac{\dot{\delta}_i}{\dot{\delta}^*}\right) + b \ln\left(\frac{\theta_i \dot{\delta}^*}{D_c}\right) \right]$$

- State evolution: $\dot{\theta}_i = 1 \frac{\theta_i \delta_i}{D}$
- Applied stress evolution:

$$\begin{split} \dot{\theta}_i &= 1 - \frac{\theta_i \delta_i}{D_c} - \frac{\alpha \theta_i}{b \sigma_i} \dot{\sigma}_i \\ \dot{\tau}_i &= \dot{\tau}_i^{\text{tect}} + K_{ij}^{\tau} \dot{\delta}_j \\ \dot{\sigma}_i &= \dot{\sigma}_i^{\text{tect}} + K_{ij}^{\sigma} \dot{\delta}_j \end{split}$$

- Terms in red are additional ones due to normal stress variations (Linker and Dieterich, 1992)
- Interaction coefficients, K, calculated from the dislocation solutions of Okada, 1992
- Tectonic stressing rates derived from backslipping the model
- Numerical integration too slow for the scale of problems we would like to address

Coulomb stress change from unit slip on one element



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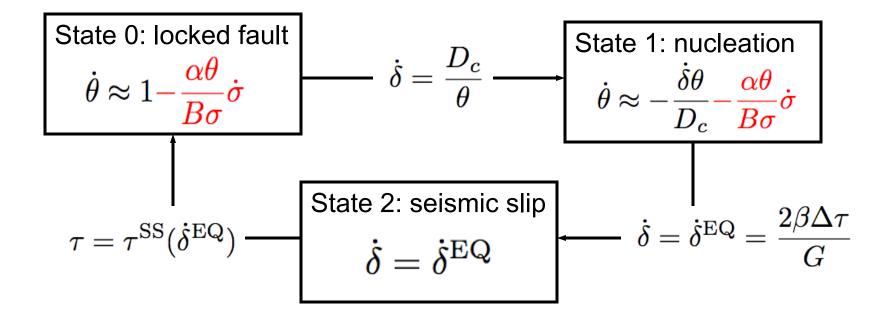
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• State evolution: $\dot{\theta}_i = 1 - \frac{\theta_i \dot{\delta}_i}{D_c} - \frac{\alpha \theta_i}{b \sigma_i} \dot{\sigma}_i$

• Stress evolution:
$$\dot{\tau}_i = \dot{\tau}_i^{\text{tect}} + K_{ij}^{\tau}\dot{\delta}_j$$
 $\dot{\sigma}_i = \dot{\sigma}_i^{\text{tect}} + K_{ij}^{\sigma}\dot{\delta}_j$



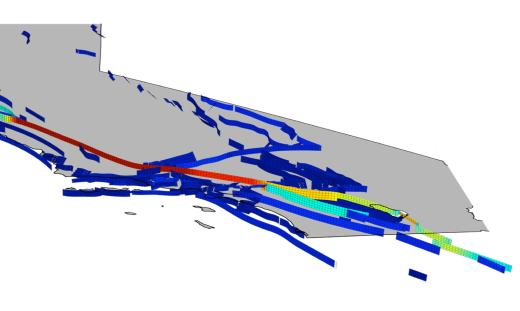
M=7.0 Multi-fault earthquake rupture simulation

Fault 2

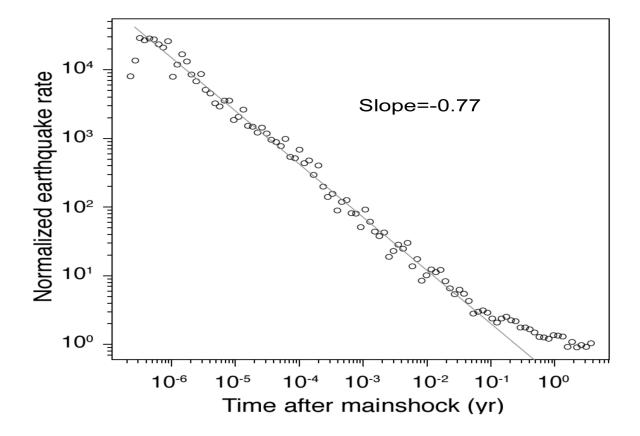
Fault 1 Red shows areas that are actively slipping in the earthquake Earth Surface Fault 2 Fault 1

Developed by Jim Dieterich and Keith Richards-Dinger at UC Riverside

- Comprehensive simulation of fault slip phenomena:
 - → earthquakes, continuous creep, slow slip events, afterslip

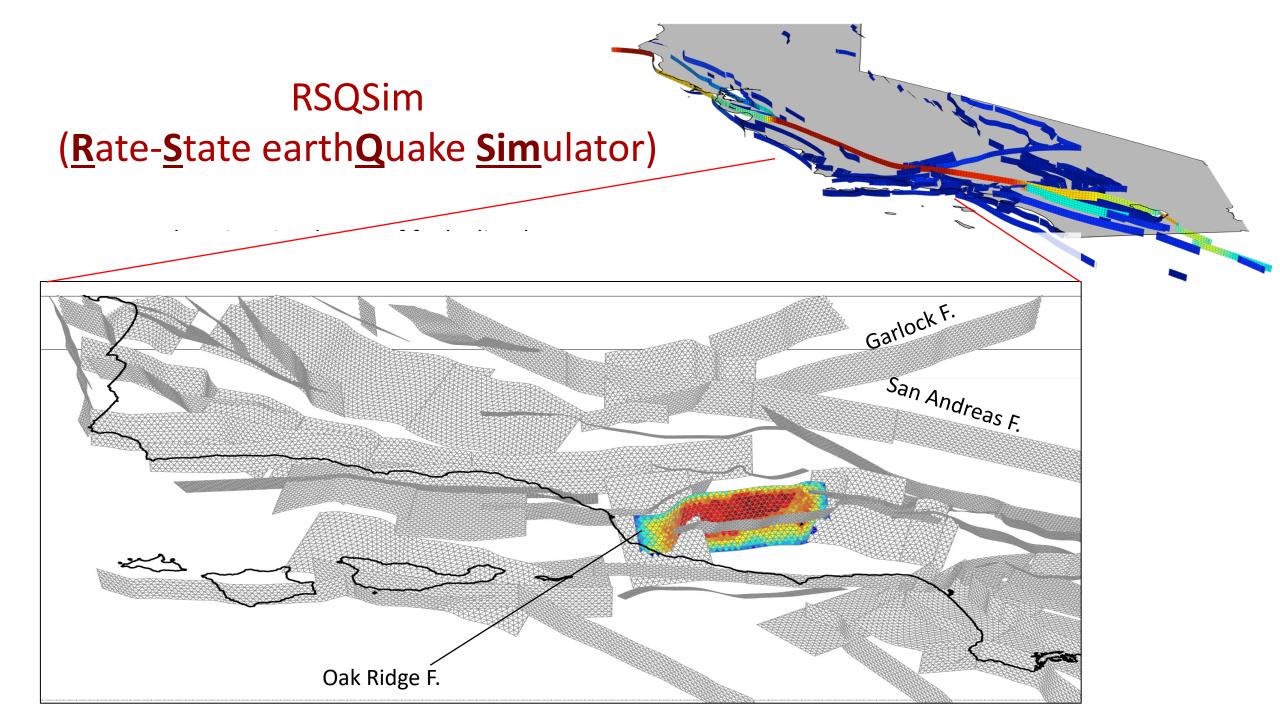


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 - → Earthquake clustering effects (aftershocks and foreshocks)



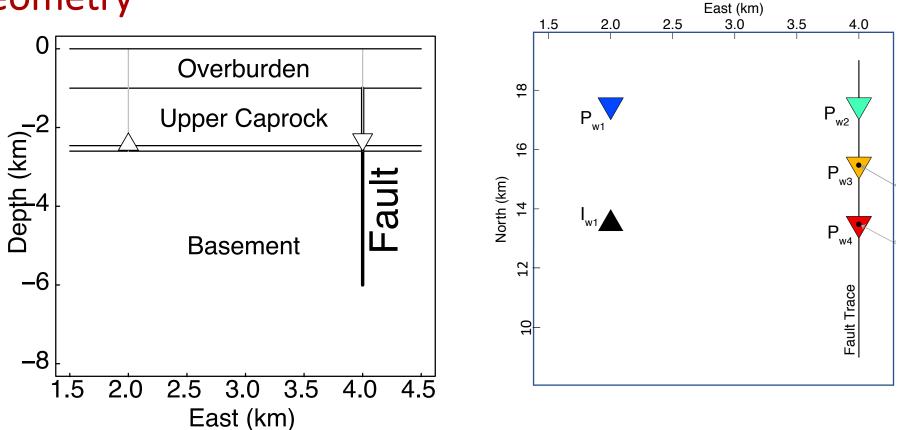
All-California simulation Aftershocks follow the Omori Law for aftershock decay with time

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- High resolution models of geometrically complex fault systems
 - \rightarrow Up to 10⁶ fault elements
 - → Range of earthquake magnitudes M=3.5 to M=8 (for 1 km² triangular elements)



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- High resolution models of geometrically complex fault systems
 - \rightarrow Up to 10⁶ fault elements
 - → Range of earthquake magnitudes M=3.5 to M=8 (for 1 km² triangular elements)
- Highly efficient code
 - → Good statistical characterizations from long simulations of 10⁶ earthquakes
 - → Repeated simulations to explore parameter space

Model Geometry



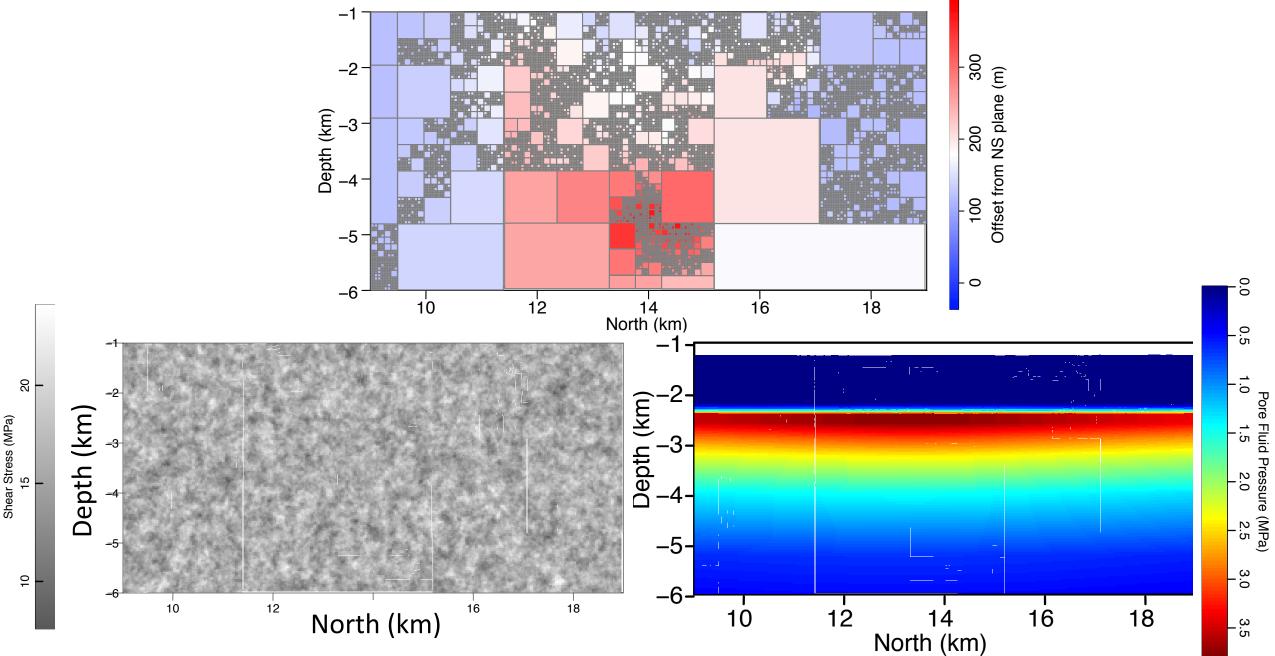
Injection and Co-production (various volumes)

- 25 % (14.25 Mt CO2)
- 50% (28.50 Mt CO2)
- 75% (42.75 MT CO2)
- 100% (57 Mt CO2)

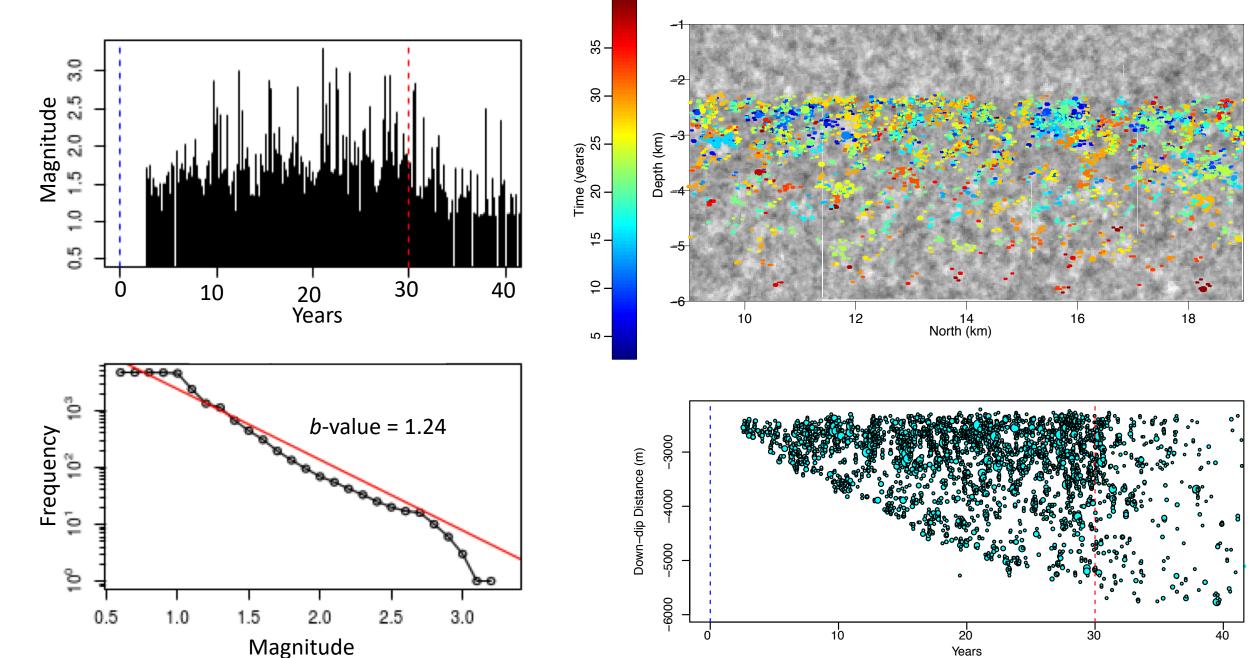
- Injection for 30 years (~0.6 m³/s)
- Co-production
 - 30 years at fractional rate

*Net injection volume = injected volume – production volume

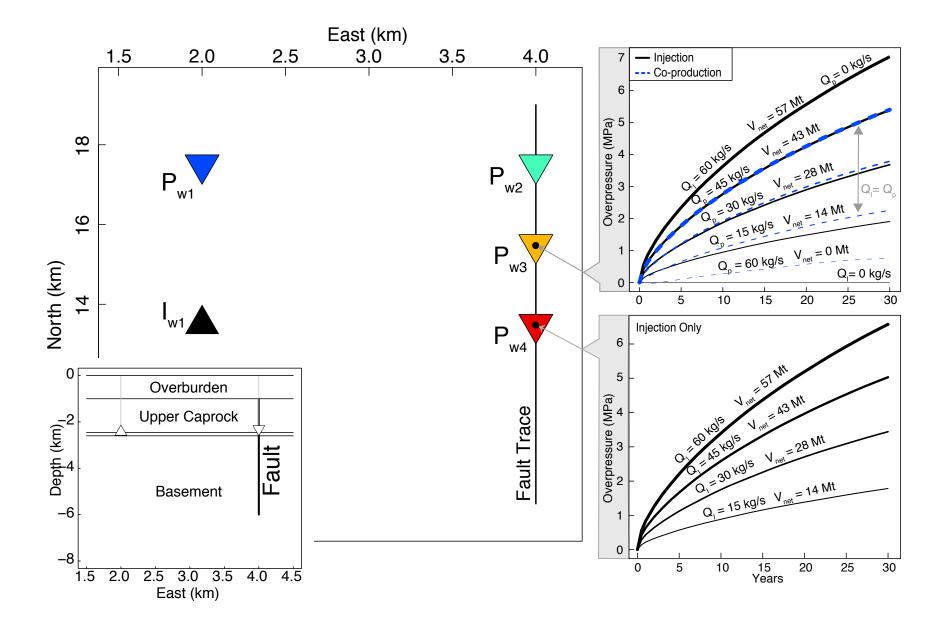
Fault Geometry, Pre-stress, and Pressure

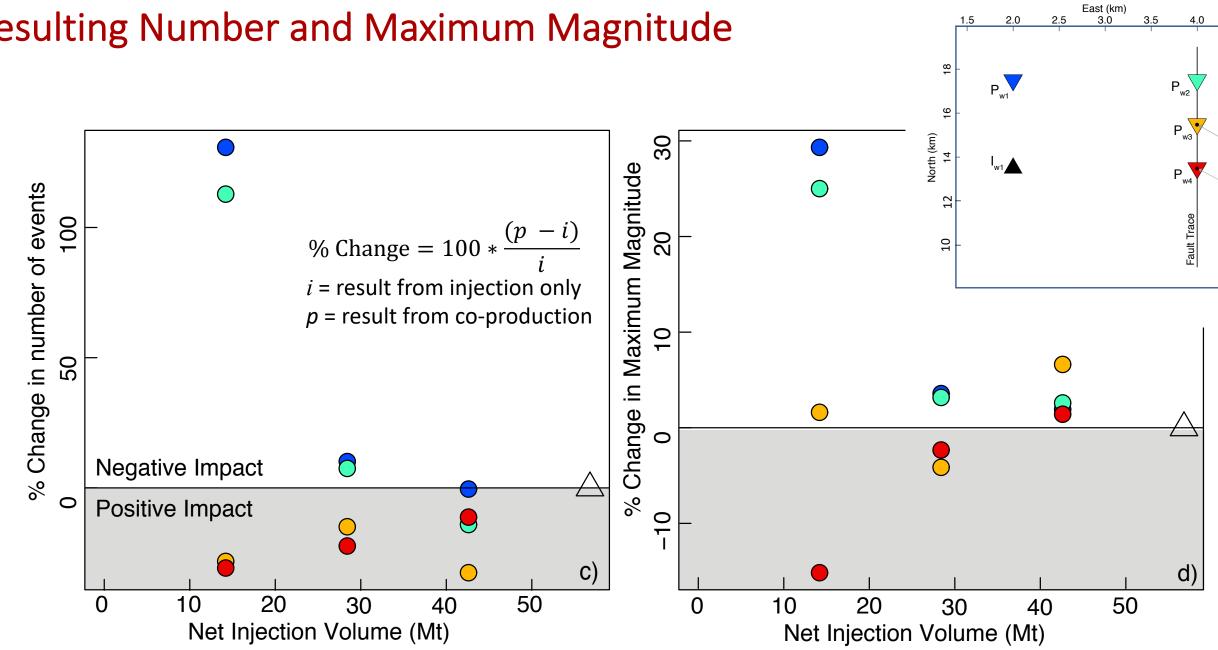


Resulting Seismic Catalog (Injection Only)



Pore-fluid Pressure On Fault



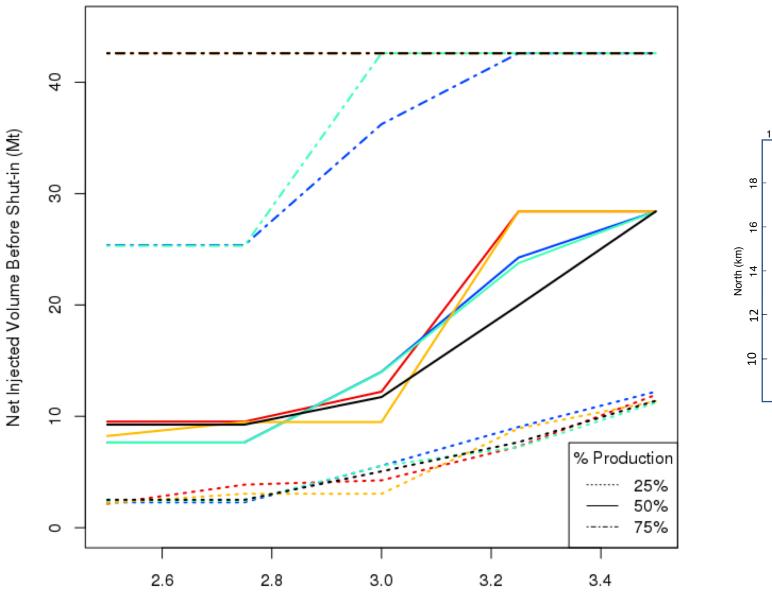


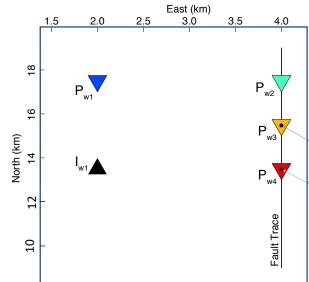
*Net injection volume = injected volume – production volume

2.0

Resulting Number and Maximum Magnitude

Selection of Preferred Injection Scenario?





Magnitude Threshold

Conclusions

- Active pressure management *might* be a useful mitigation strategy
 - May actually cause more (but smaller?) earthquakes
 - Highly dependent upon knowledge of fault location, well/fault configuration, reservoir characteristics, the pre-stress conditions, fault interaction and permeability structure
- Most advantageous when co-produced fluids can be managed and potentially dangerous faults are known
- Perhaps more applicable to carbon storage settings?
 - Replace low-risk brine with high-risk CO2
 - Fewer coordination logistics