

# Parametrization Technique for Reactive Multiphase Flow and Transport

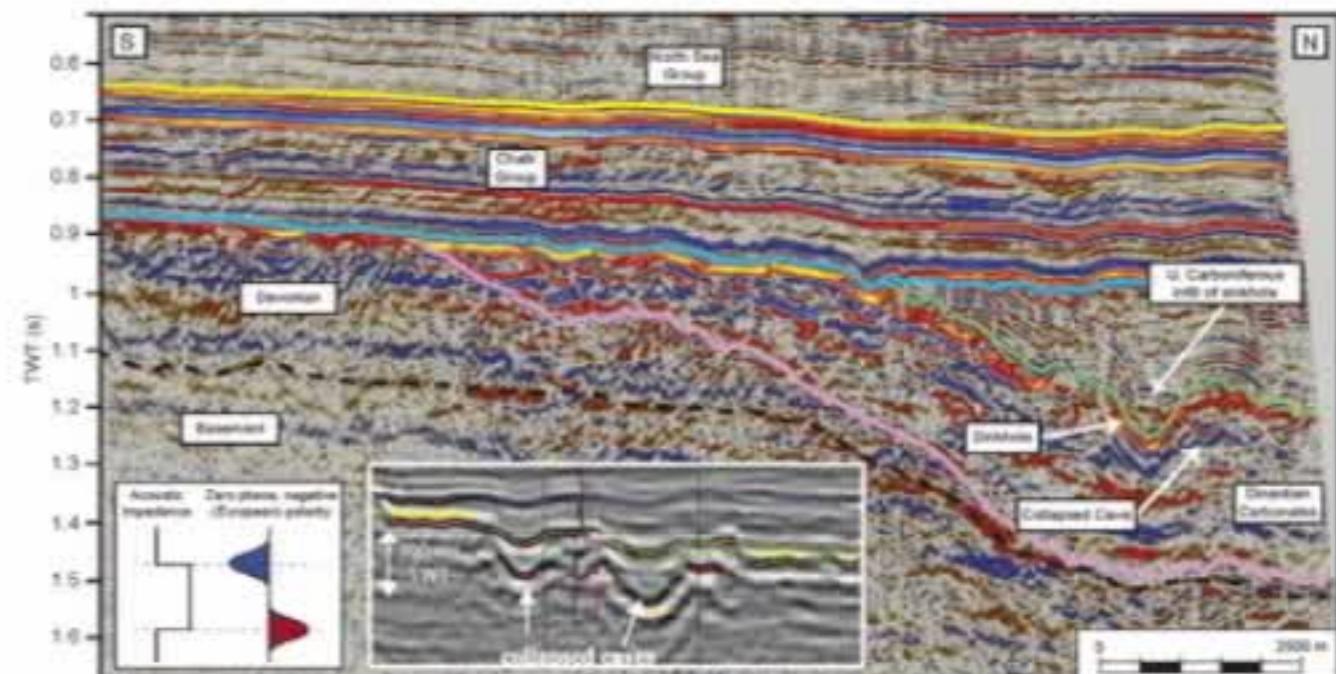
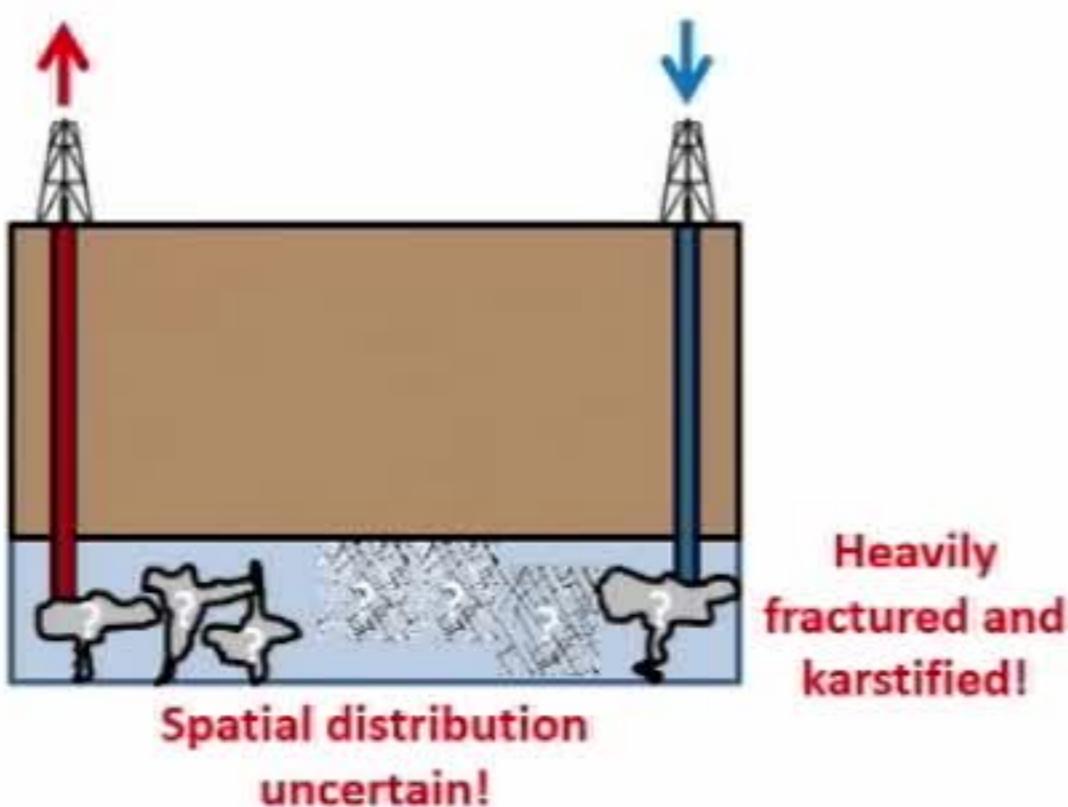
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# Motivation

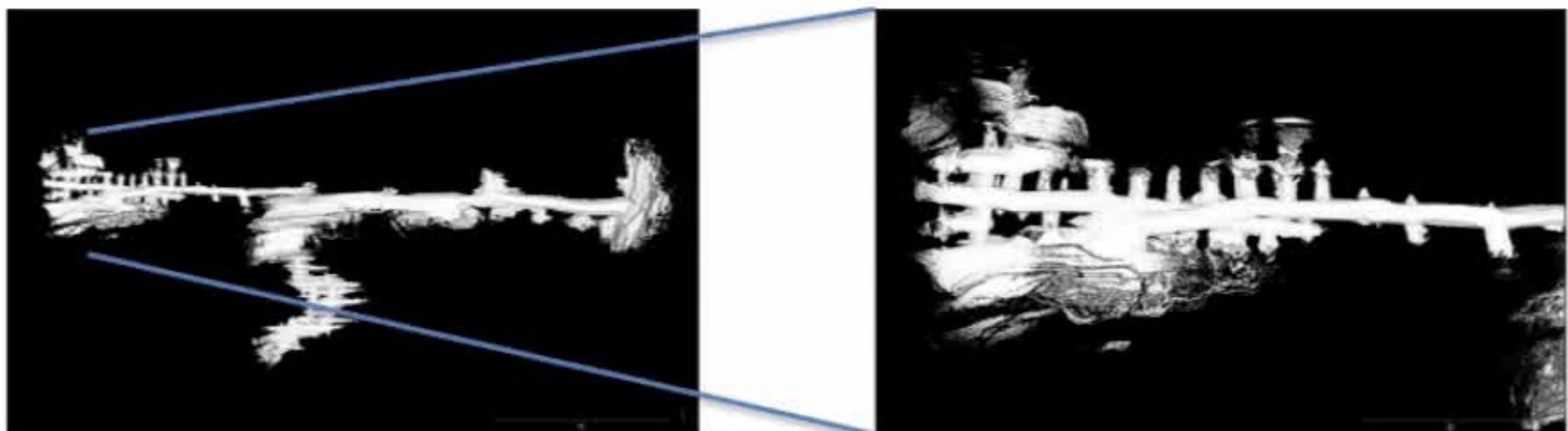
- **Hydrocarbon and Geothermal** energy potential in (pre-salt) carboniferous carbonates
- Difficult reservoir characterization, large uncertainty in spatial distribution of discontinuity networks → **unpredicted hazards**



Reijmer et al. (2017)

# Objective

Reduce risks of subsurface applications by improving the  
geomodeling workflow of fractured & karstified reservoirs.  
This requires integration of coupled flow and reactive  
transport in porous media with outcrop and lab data.



# Linearization

- Implementation depends on a
  - nonlinear formulation
  - number of phases and components
  - fluid and rock properties
  - etc...
- Involves huge volumes of code

The most difficult part to design and optimize  
(especially for new hardware architectures)

## Which data involved in simulation?



SGOF			
0	0	1	0
0.001	0	1	0
0.02	0	0.997	0
0.05	0.005	0.98	0
0.12	0.025	0.7	0
0.2	0.075	0.35	0
0.25	0.125	0.2	0
0.3	0.19	0.09	0
0.4	0.41	0.021	0
0.45	0.6	0.01	0
0.5	0.72	0.001	0
0.6	0.87	0.0001	0
0.7	0.94	0	0
0.88	0.98	0	0

$$\beta_c(\omega) = \sum_{j=1}^{n_p} x_{cj}^l \rho_j^l \frac{k_{rj}^l}{\mu_j^l}.$$

$$p = \frac{RT}{v - b} - \frac{a}{v(v + b) + b(v - b)}$$

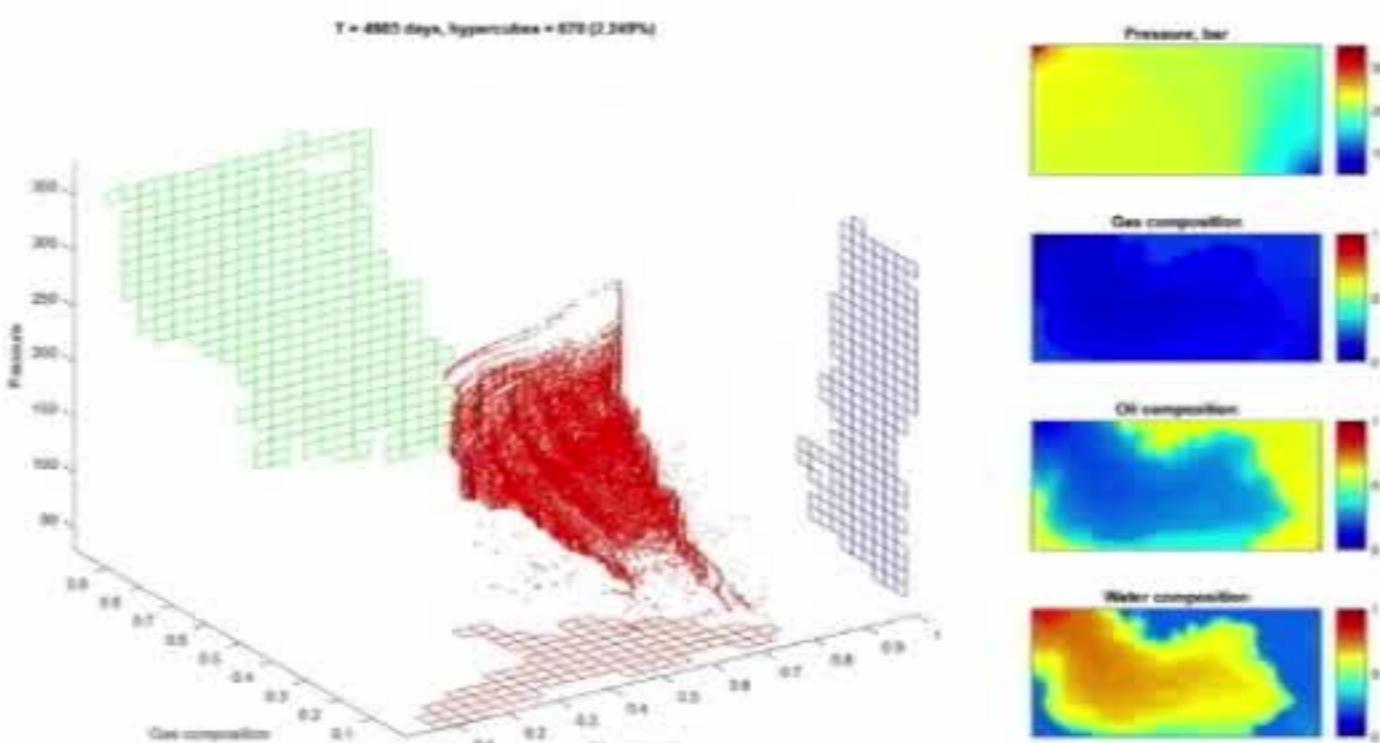
$$\left[ (\mu - \mu^0) \xi_T + 10^{-4} \right]^{1/4} = a_0 + a_1 \rho_{pr} + a_2 \rho_{pr}^2 + a_3 \rho_{pr}^3 + a_4 \rho_{pr}^4$$



# Simulation framework for inversion

## Delft Advanced Research Terra Simulator

- Discretization in space, time and **physics** (OBL)
- C++/CUDA implementation of performance-critical kernels + **open-source** Python interfaces
- Validated for various geothermal and petroleum applications
- Adaptive parametrization in space of unknowns



More information at <https://darts.citg.tudelft.nl/>

# Compositional simulation with reactions

Reducing number of mass balance equations:

- Aris and Mahi, 1963: Reaction invariants
- Rubin, 1983: Tenads
- Lichtner, 1985 & 1996: Primary species
- Molins et al., 2004: Components
- Fan et al., 2012: Elements

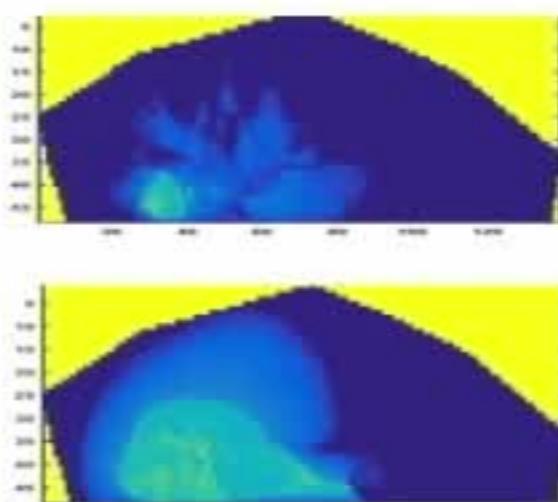
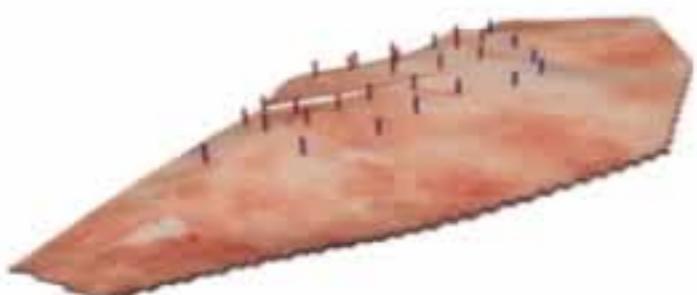
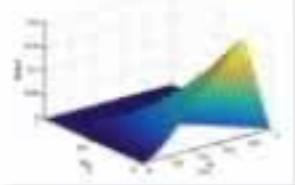
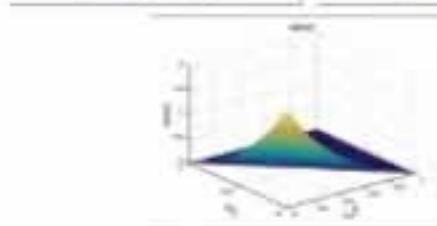
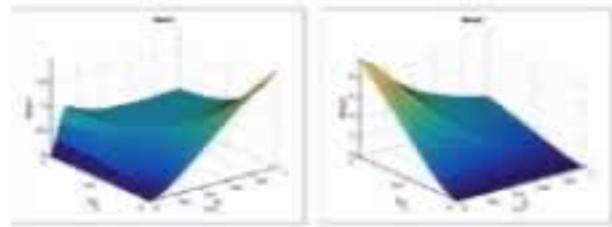
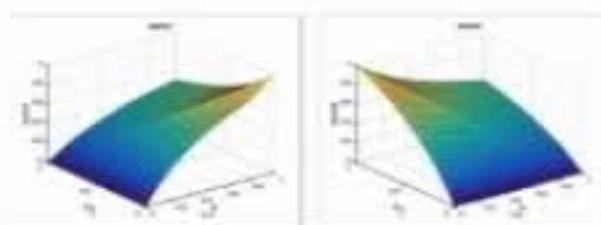
Multiphase thermal multicomponent reactive transport  
(Sequential or Fully Implicit schemes):

- TOUGHREACT (Xu et al., 2012)
- PFLOTRAN (Lichtner et al., 2015)
- DUMUX (Flemisch et al., 2011)
- AD-GPRS (Farshidi, 2016)

# Equilibrium reactions with OBL

$$E \times \frac{\partial}{\partial t} (\phi \rho_t z_c) + \operatorname{div}(l_c) = \sum_{q=1}^{n_q} v_{cq} r_q \quad \rightarrow \quad \frac{\partial}{\partial t} (\phi^T \rho_t^E z_i^E) + \operatorname{div}(\mathbf{e}_i l) = 0$$

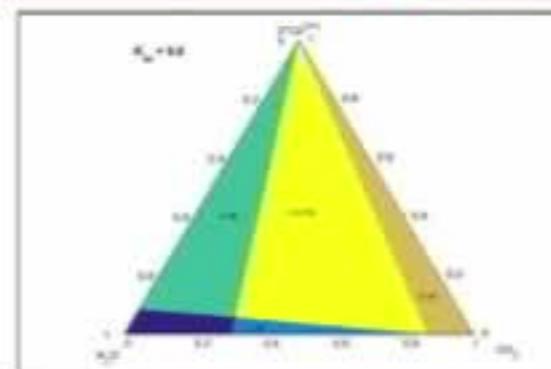
$$\boldsymbol{\omega} = \{p, \mathbf{z}^E\} \rightarrow \begin{cases} \alpha_i(\boldsymbol{\omega}) = c(p) \rho_t^E z_i^E \\ \beta_i(\boldsymbol{\omega}) = \mathbf{e}_i \sum_{j=1}^{n_p} \mathbf{x}_j \rho_j \frac{k_{rj}}{\mu_j} \end{cases}$$



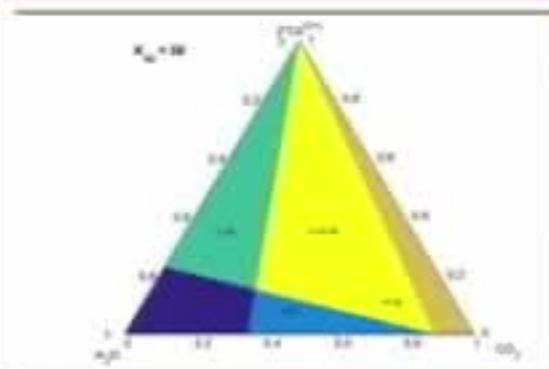
$$\prod_{c=1}^{n_c} a_c^{v_{cq}} - K_q = 0$$

$$\rho_t^E = \rho_t \sum_{i=1}^{n_e} \mathbf{e}_i \mathbf{z}$$

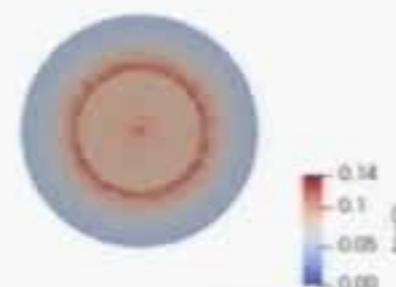
$$\mathbf{z}^E \sum_{i=1}^{n_e} \mathbf{e}_i \mathbf{z} - E \mathbf{z} = 0$$



Grid A



Grid B



Grid D

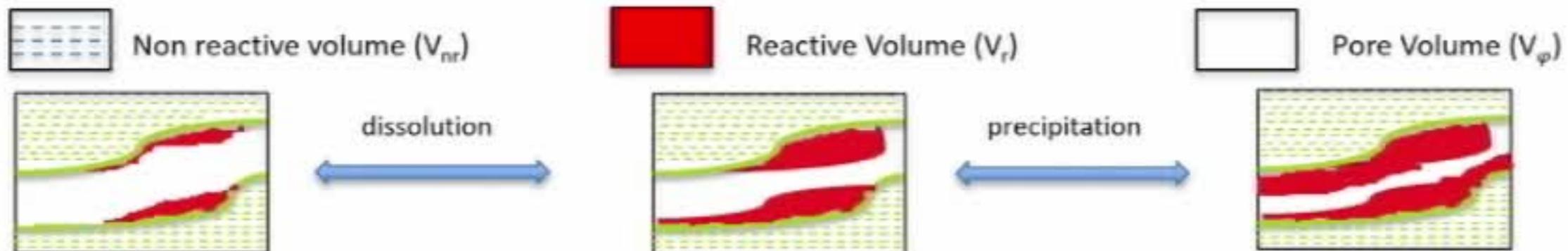


Grid C



Grid D

# Porosity-permeability treatment



- Bulk volume is a sum of reactive ( $V_r$ ), non reactive ( $V_{nr}$ ) and pore volume ( $V_\varphi$ ).

$$V_b = V_\varphi + V_r + V_{nr}$$

- We introduce the concept of total porosity which is defined as

$$\varphi^T = \varphi_r + \varphi$$

- Fluid porosity can be calculated using the mineral saturation values

$$\varphi^{n+1} = \varphi^T \left(1 - \sum_i^{n_m} S_i^{n+1}\right)$$

- Power law transmissibility multiplier using upstream fluid porosity

$$T^M(\varphi^{n+1}) = \left(\frac{\varphi_{n+1}}{\varphi^T}\right)^A = \left(1 - \sum_i^{n_m} S_i^{n+1}\right)^A$$

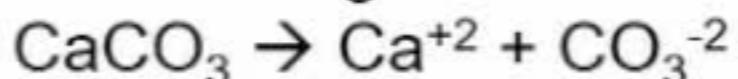
- In this way,  $T^M \in [0, 1]$ , i.e. at any time:  $T = \Gamma^{max} T^M$

# Operator form of equations

$$a(\xi)(\alpha_c(\omega) - \alpha_c(\omega_n)) + \sum_l b(\xi, \omega) \beta_c(\omega) - c(\xi) \gamma_c(\omega) = 0$$

- $a(\xi) = \phi_0(\xi)V(\xi), \quad \left| \begin{array}{l} \alpha_c(\omega) = (1 + c_r(p - p_{ref})) \sum_{j=1}^{n_p} x_{cj} \rho_j s_j, \\ \beta_c(\omega) = \sum_{j=1}^{n_p} T^M x_{cj}^l \rho_j^l \frac{k_{rj}^l}{\mu_j^l}, \\ \gamma_c(\omega) = \sum_{k=1}^{n_k} v_{ck} r_k \end{array} \right.$
- $b(\xi, \omega) = \Delta t \Gamma^{l, max}(\xi)(p - p^l)$
- $c(\xi) = \Delta t V(\xi),$

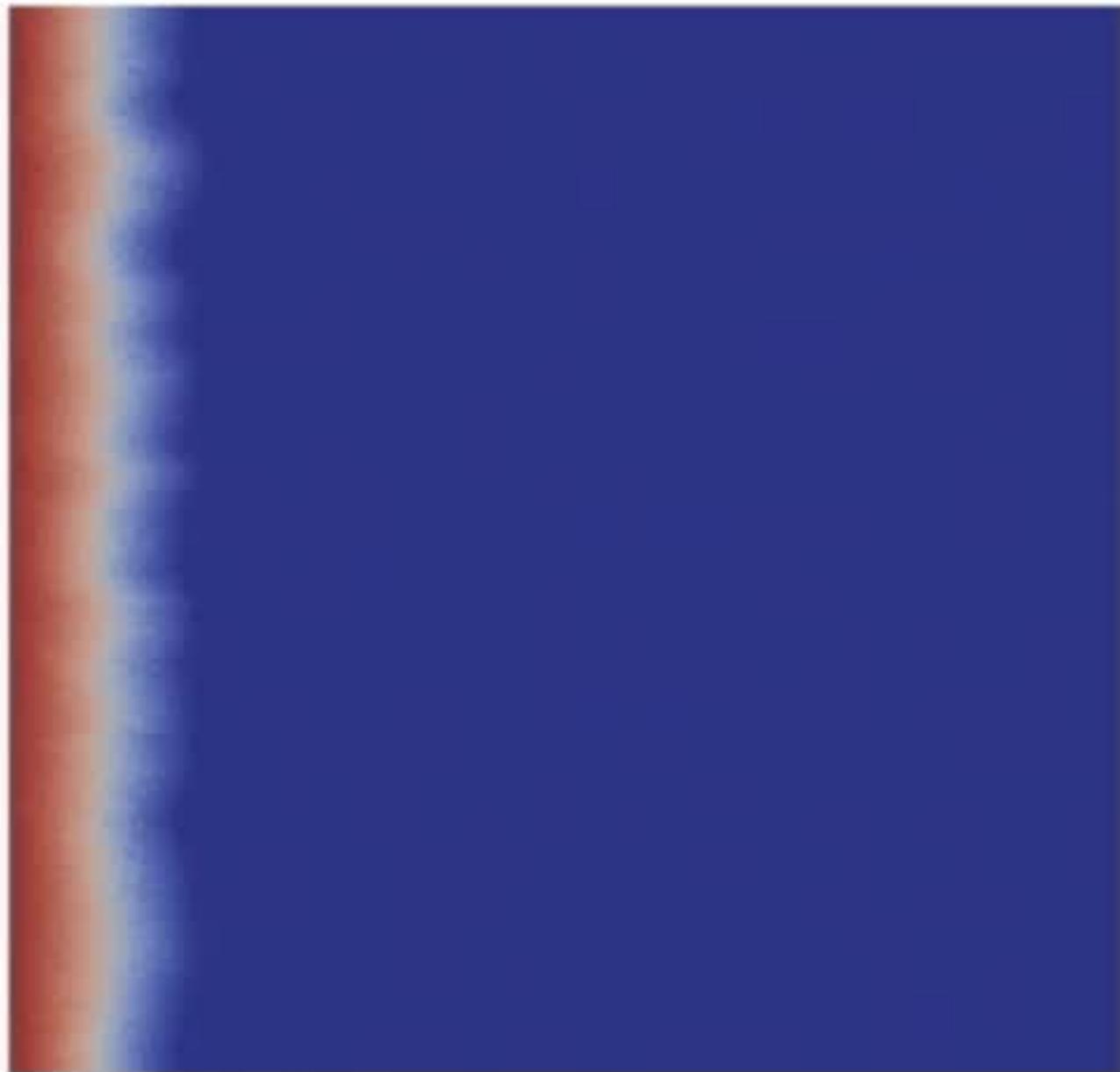
Example for a single kinetic reaction:



$$\gamma_c(\omega) = v_c k A_s \left(1 - \frac{Q}{K}\right) z_{\text{CaCO}_3}$$

# Slower kinetic rates

$$L_{res} = 100 \text{ [m]}; H_{res} = 10 \text{ [m]}; Q_{inj} = 1 \text{ [m}^3/\text{day}] \rightarrow v_{avg} = 0.001 \text{ [m/day]}$$



$$k_{rate} = 0.0005 \text{ [1/day]} \rightarrow Da \sim 0.5$$



$$k_{rate} = 0.00001 \text{ [1/day]} \rightarrow Da \sim 0.01$$

## Conclusion and future work

- Parametrization approach extended for kinetic and equilibrium reactions
- Flexibility of the OBL framework helps in implementation of complex physics

### On-going work

- Validate simulation results with reactions
  - CO<sub>2</sub> sequestration
  - Gas-hydrates
  - Geothermal scaling
  - Chemical EOR
- Add diffusion operators
- Apply framework to fracture network
- Link to real cave data and laboratory experiments

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## Conclusion and future work

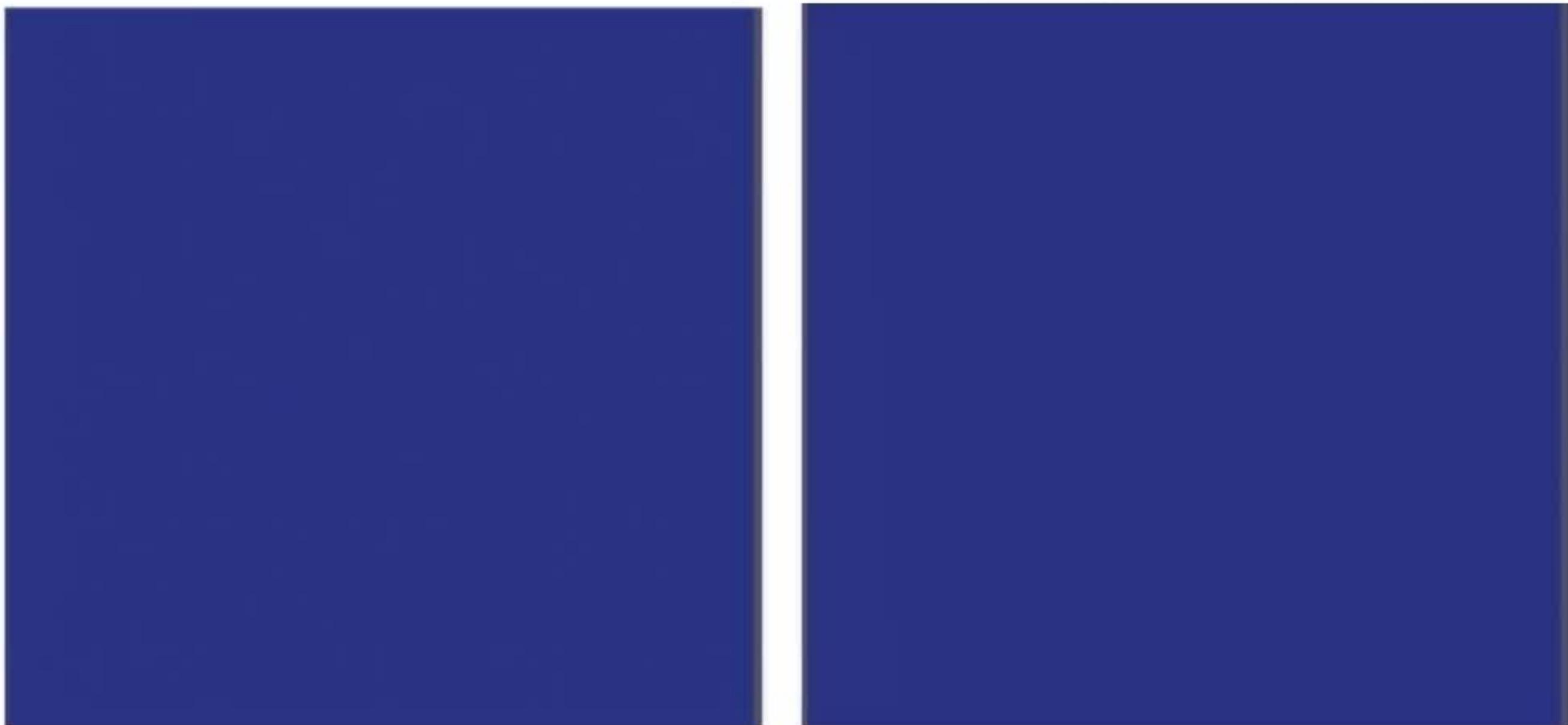
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# Wormholes in different discretizations

$L_{res} = 100 \text{ [m]}$ ;  $H_{res} = 10 \text{ [m]}$ ;  $Q_{inj} = 1 \text{ [m}^3/\text{day}] \rightarrow v_{avg} = 0.001 \text{ [m/day]}$



## Kinetic reactions with OBL

$$\frac{\partial}{\partial t} (\varphi^T \rho^T z_c) + \operatorname{div} \sum_{j=1}^{n_p} x_{cj} \rho_j \vec{u}_j - \sum_{k=1}^{n_k} v_{ck} r_k - \sum_{q=1}^{n_q} v_{cq} r_q = 0$$

$$(V \varphi^T \rho^T z_c)^{n+1} - (V \varphi^T \rho^T z_c)^n - \Delta t \sum_{l \in L} \sum \left( x_{cj} \rho_j \lambda_j \Gamma^{\max} T^M \right)^l (p - p^l) \\ - \Delta t \left( V \sum_{k=1}^{n_k} v_{ck} r_k \right)^{n+1} = 0$$

$$a(\xi)(\alpha_c(\omega) - \alpha_c(\omega_n)) + \sum_l b(\xi, \omega) \beta_c(\omega) - c(\xi) \gamma_c(\omega) = 0$$