

# Ephemeral Transport Boundaries in Geophysical Flows

H. Chang

H. S. Huntley

A. D. Kirwan, Jr.



# Thanks To



CONSORTIUM for ADVANCED  
RESEARCH and TRANSPORT of  
HYDROCARBONS in the  
ENVIRONMENT



Support from Office of Naval Research through  
MURI *OCEAN 3D + 1* grant N00014-11-1-0087



Professor Daniel Karrasch  
Professor George Haller



# Principle of Objectivity

- Concept applied by Noll (1955, 1958, 1959) to viscous processes in non-Newtonian fluids
- Speziale et al. (1981, 1991, 1996) applied concept to turbulence closure schemes
- Haller et al. (2005, 2013, 2015, 2016, 2017) applied concept to identification of LCS

*Unifying theme: Reconcile theory and observations for nonlinear behavior of fluids*

# Exordium

*Material properties are independent of the observer. – W. Noll*

*As far as the laws of mathematics refer to reality, they are uncertain, and as far as they are uncertain, they do not refer to reality. – A. Einstein*

Some results are just days old!

# Goals

1. Oceanographers should worry about Objectivity. I shall illustrate why.
2. Study the impact of traveling internal waves on mesoscale transport pathways.

# Procedure

Flow fields from solutions to linear, stratified, 3D Euler equations on f-plane.

---

# Stratified 3D Euler Equations on f-plane

---

# Euler Equations

$$\frac{\partial u}{\partial t} - fv = -\rho_0^{-1} \frac{\partial p}{\partial x}$$

$$\frac{\partial v}{\partial t} + fu = -\rho_0^{-1} \frac{\partial p}{\partial y}$$

$$\frac{\partial w}{\partial t} = \rho_0^{-1} \left( \rho g - \frac{\partial p}{\partial z} \right)$$

$$\frac{\partial \rho}{\partial t} = \left( \frac{\rho_0 N^2}{g} \right) w$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$$

# Solution

$$u = -lA_g(z) \sin kx \cos ly - \sum_j \left( A_j(z) \frac{\partial \Psi_j(x, y, t)}{\partial y} - \frac{dB_j(z)}{dz} \frac{\partial \Phi_j(x, y, t)}{\partial x} \right)$$

$$v = kA_g(z) \cos kx \sin ly + \sum_j \left( A_j(z) \frac{\partial \Psi_j(x, y, t)}{\partial x} + \frac{dB_j(z)}{dz} \frac{\partial \Phi_j(x, y, t)}{\partial y} \right)$$

$$w = - \sum_j B_j(z) \nabla_h^2 \Phi_j(x, y, t) = \sum_j (k_j^2 + l_j^2) B_j(z) \Phi_j(x, y, t)$$

$$p = p_0(z) + p_g(z) \sin kx \sin ly + \sum_j p_j(z) \Psi_j(x, y, t)$$

$$\rho = \rho_0(z) + \rho_g(z) \sin kx \sin ly + \sum_j \rho_j(z) \Psi_j(x, y, t)$$

# Eigen Relations

$$p(x, y, z, t) = p_0 + p_g \sin kx \sin ly + p_j \Psi_j$$

$$\left[ \frac{d^2}{dz^2} + \left( \frac{N^2}{g} \right) \frac{d}{dz} + \left( \frac{N^2}{gh_j} \right) \right] p_j = 0$$

$$\left\{ \left[ \nabla_h^2 - \left( \frac{\partial^2}{\partial t^2} + f^2 \right) (gh_j)^{-1} \right] + \beta \mathcal{L} \right\} \Psi_j = 0$$

$$\beta = 0$$

# Dynamic Constraints

$$\omega_j^2 = f^2 + \frac{N^2 H^2}{L_j^2}$$

$$\frac{\partial \Psi_j}{\partial t} = -\omega_j \Phi_j$$

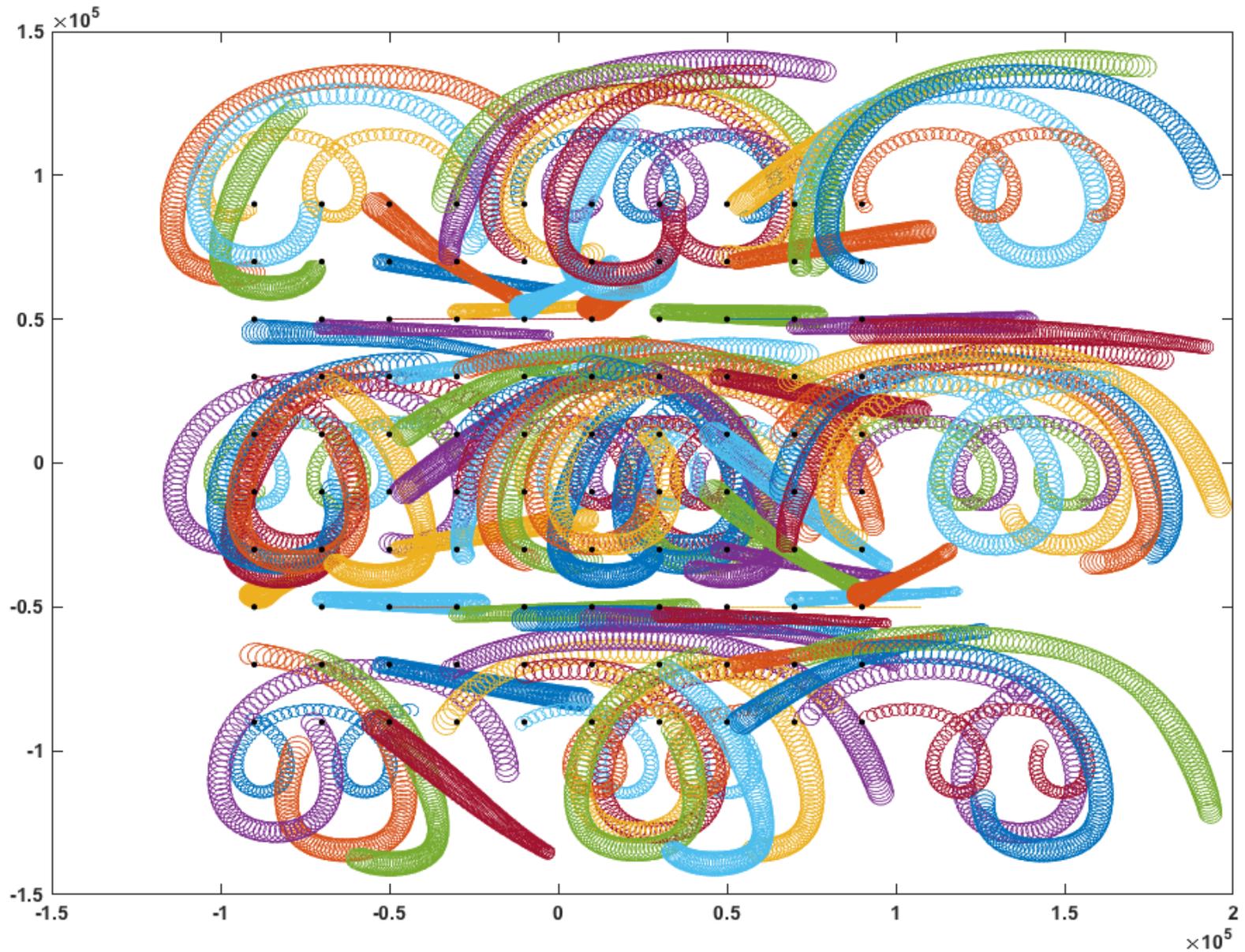
$$\frac{\partial \Phi_j}{\partial t} = \omega_j \Psi_j$$

---

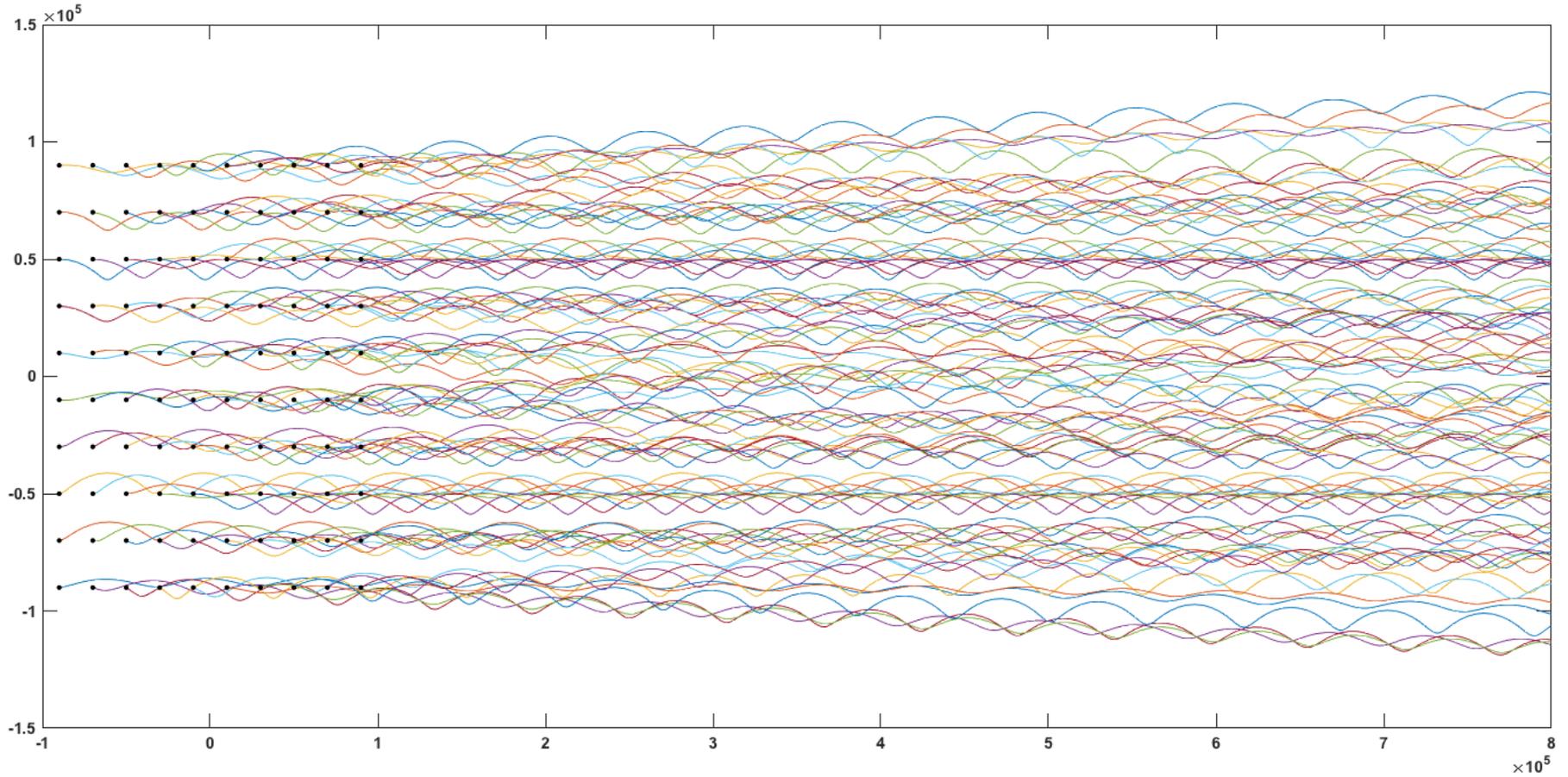
# **Objectivity – Comparison of 2 Flow Realizations**

---

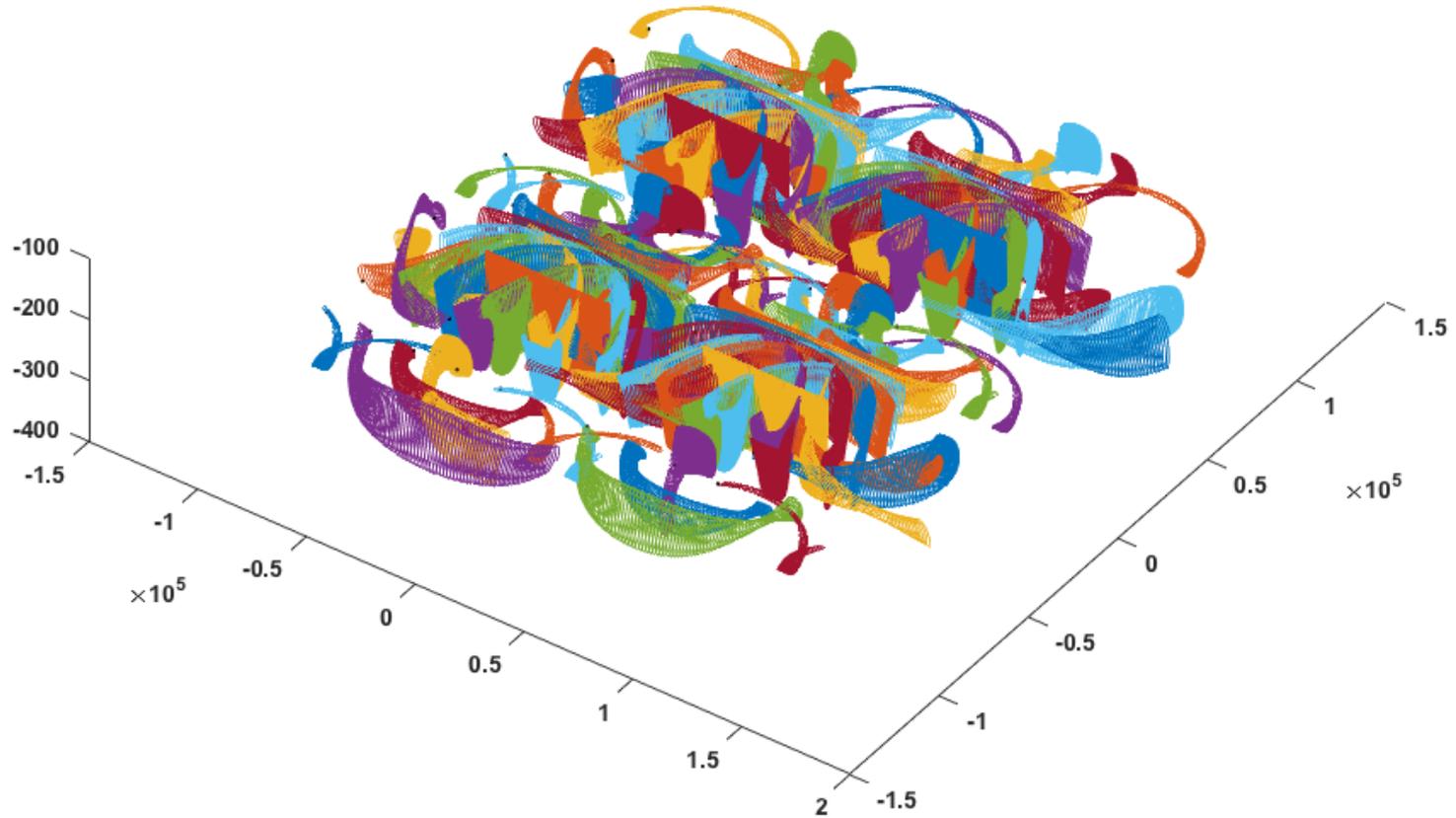
# Surface Realization 1



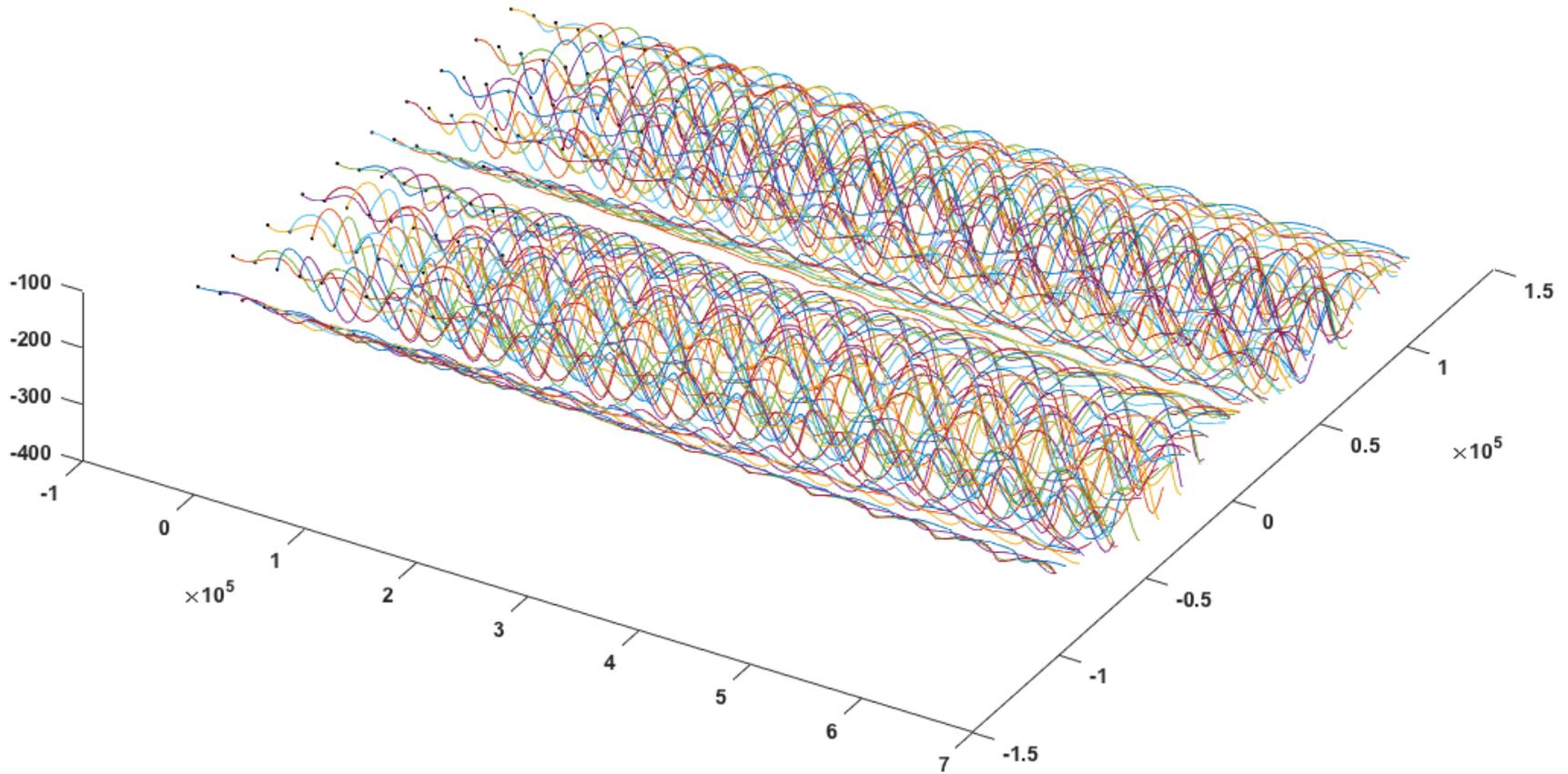
# Surface Realization 2



# Subsurface Realization 1



# Subsurface Realization 2

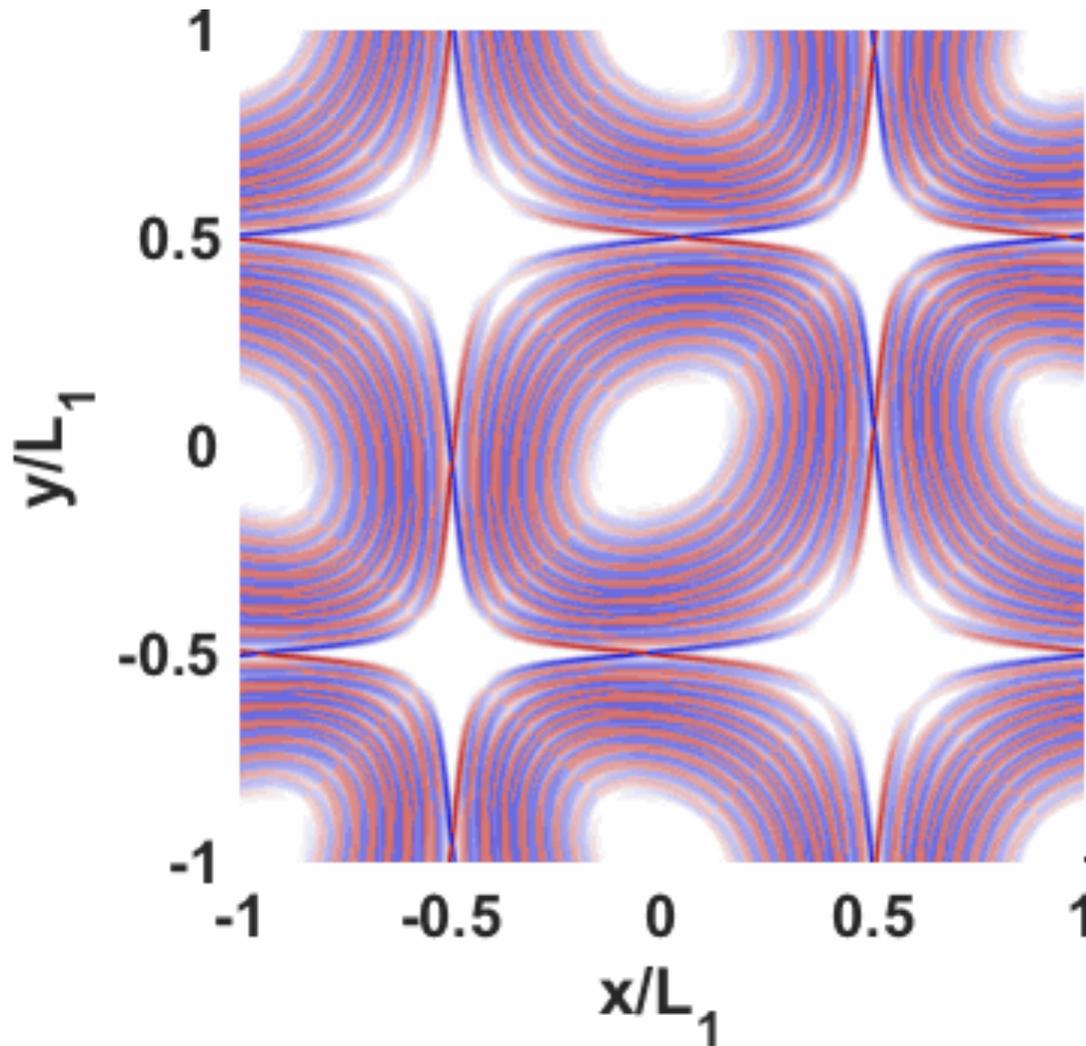


---

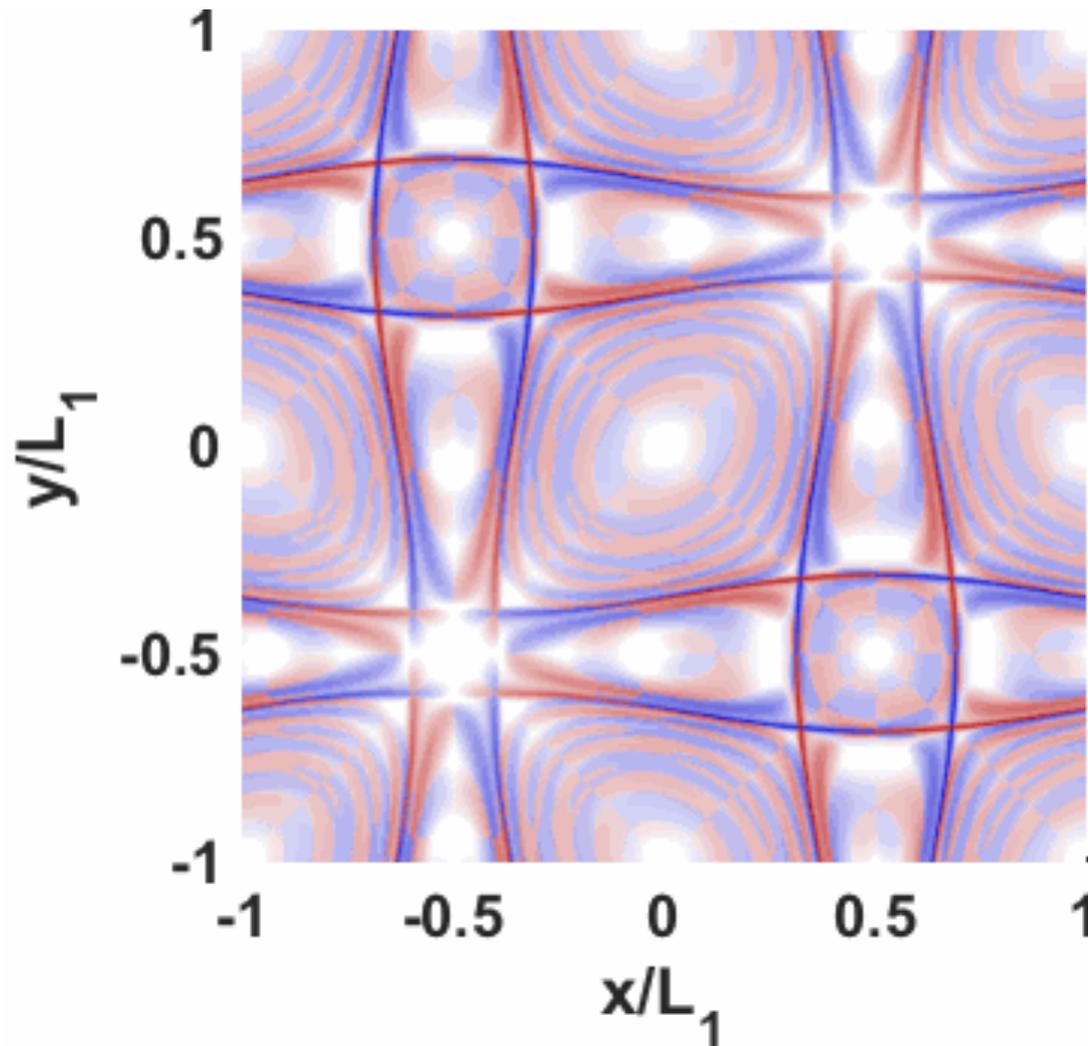
# **Objective Diagnostics – FTLE**

---

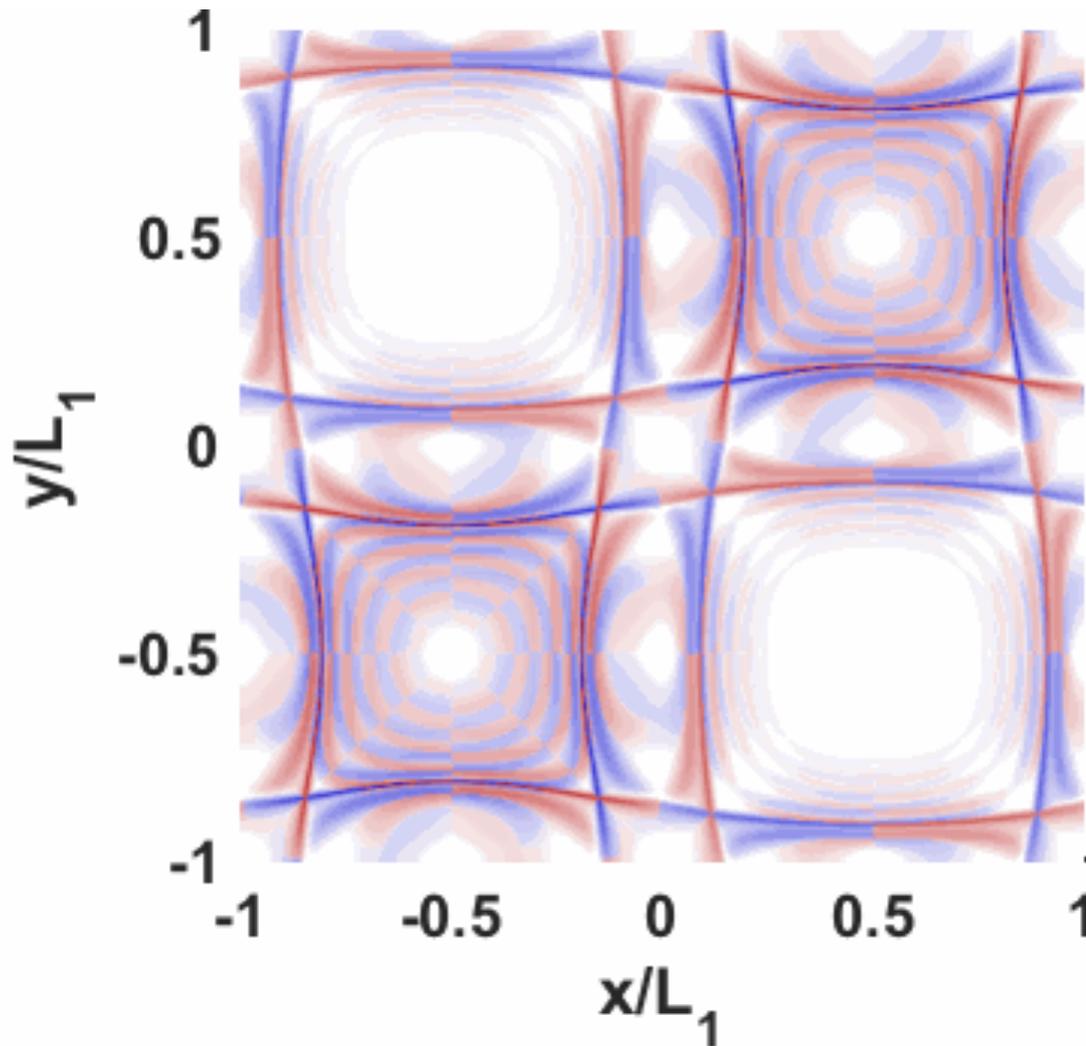
# FTLE at $z = 0$



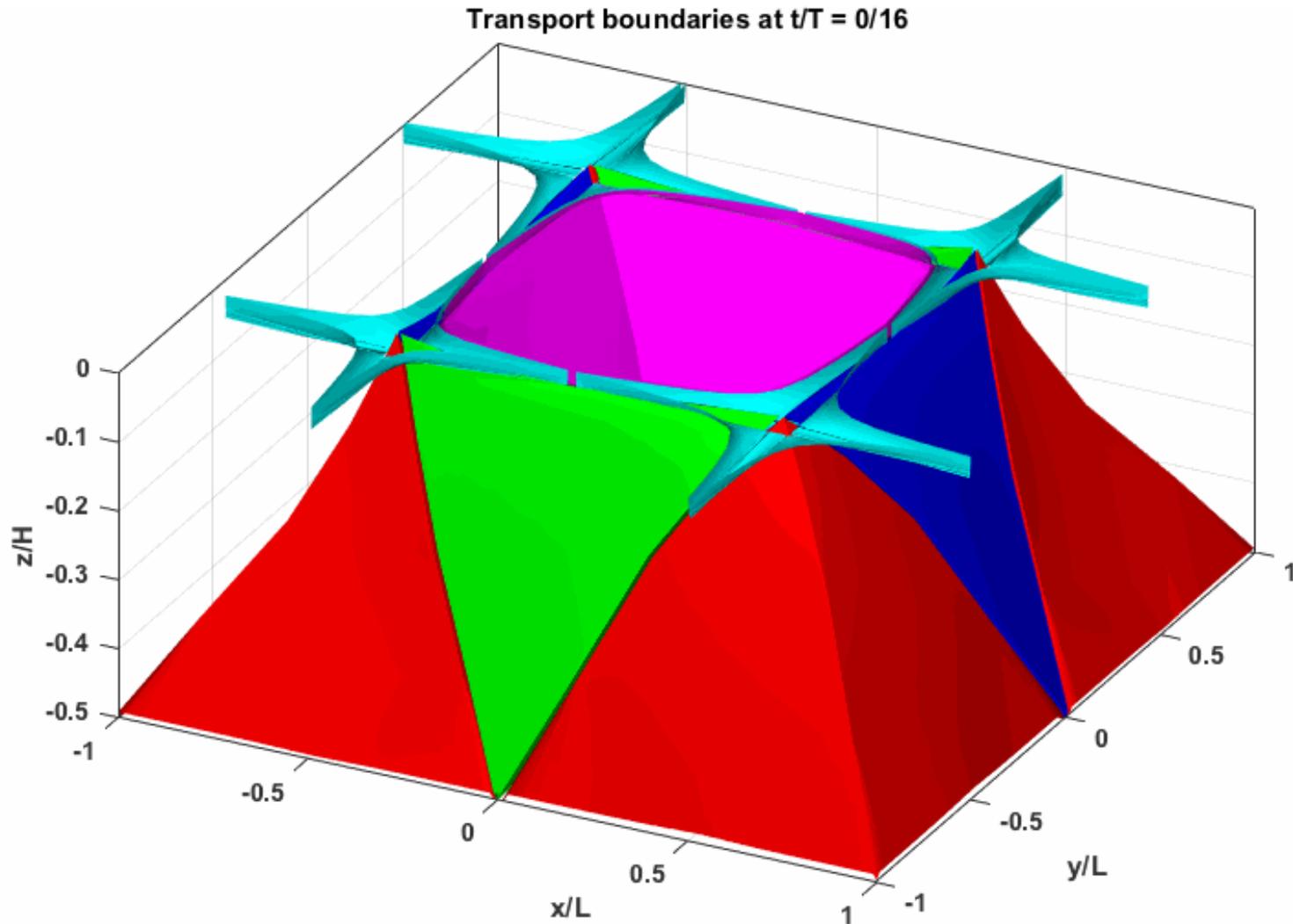
# FTLE at $z = H/8$



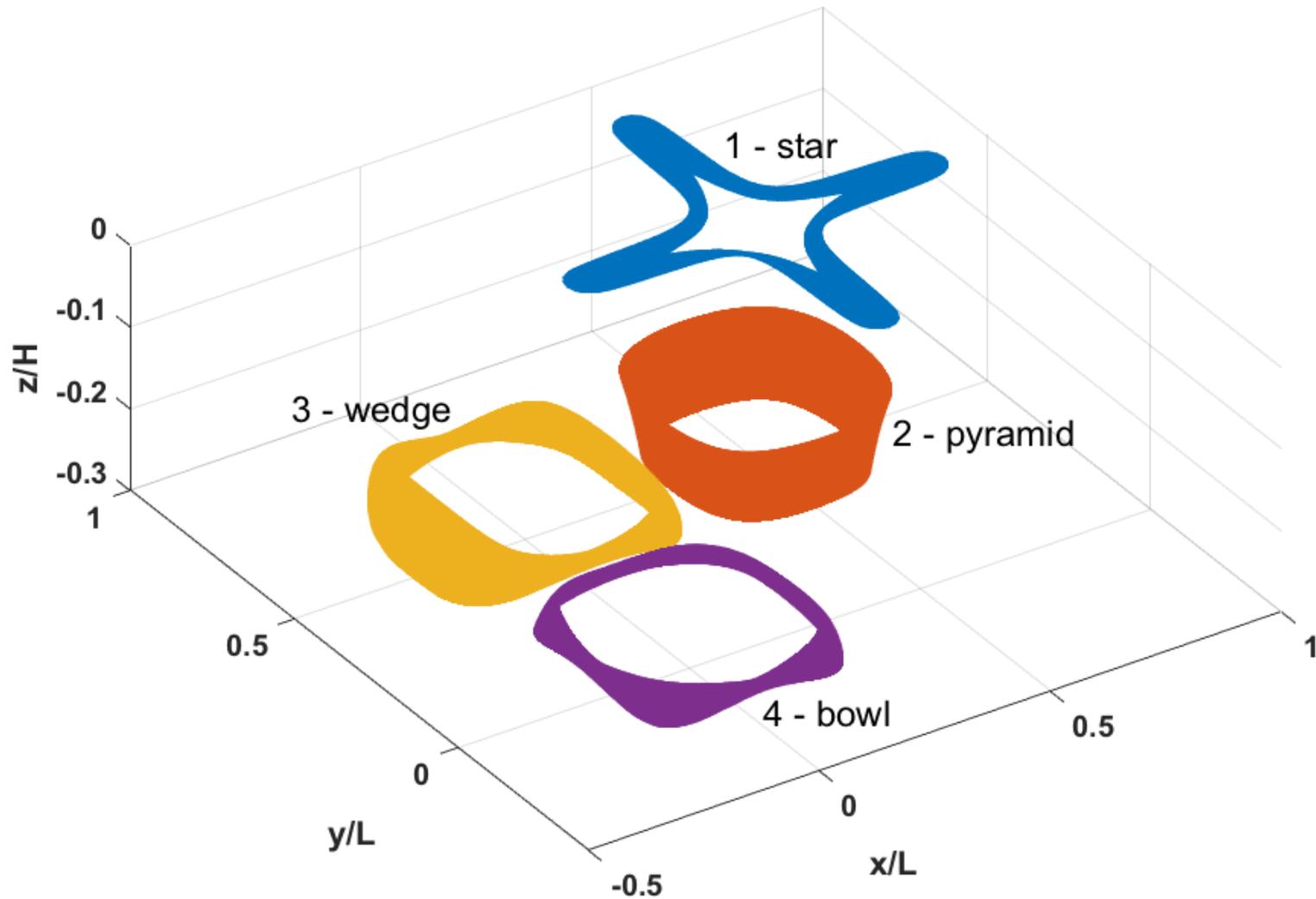
# FTLE at $z = 3 H/8$



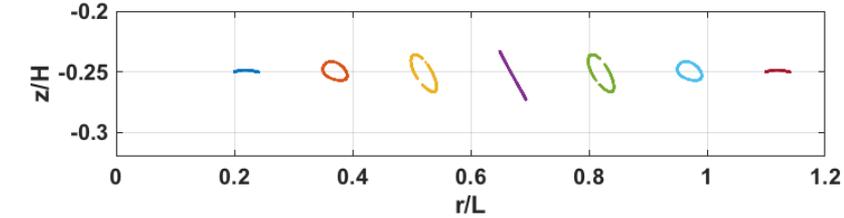
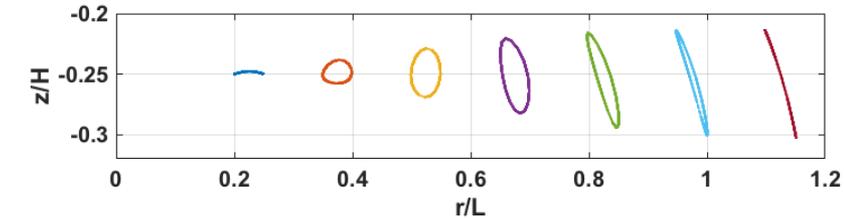
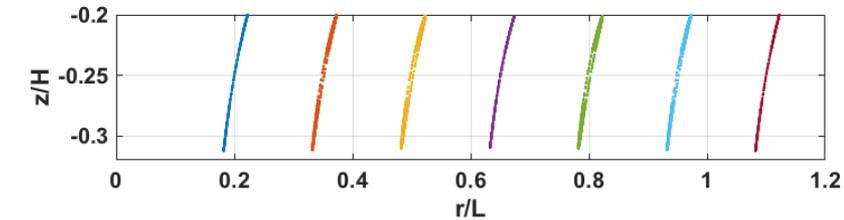
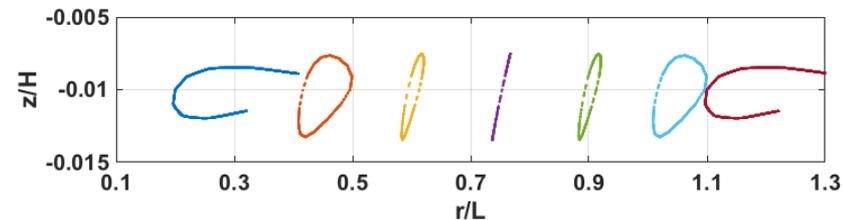
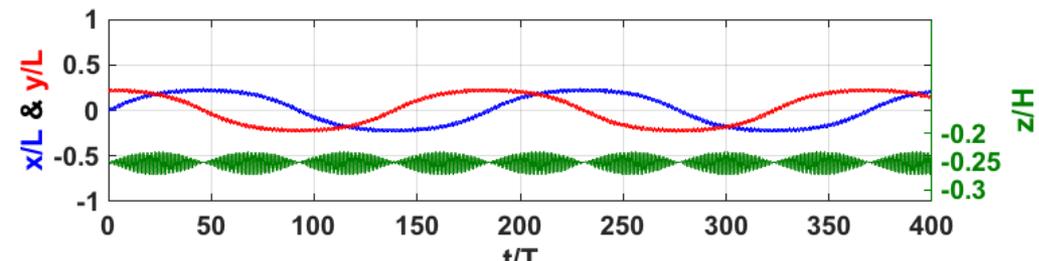
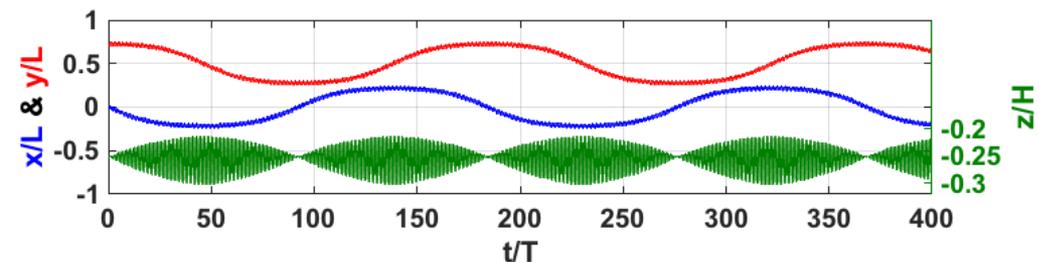
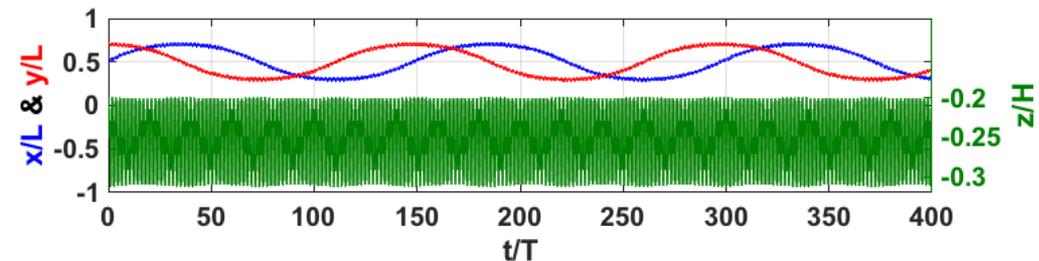
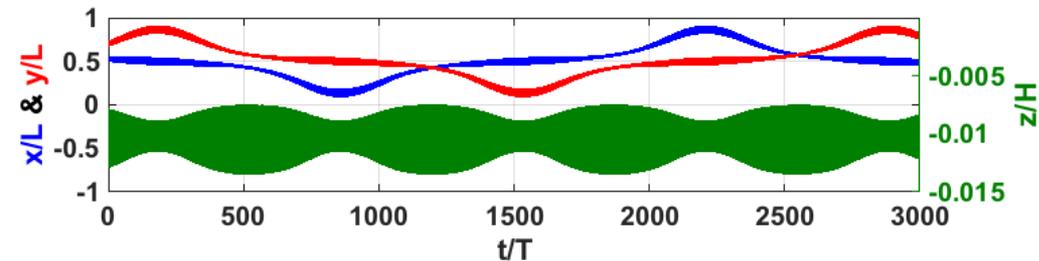
# Transport Boundaries



# Representative 3D Trajectories



# Time Series & Poincare Sections



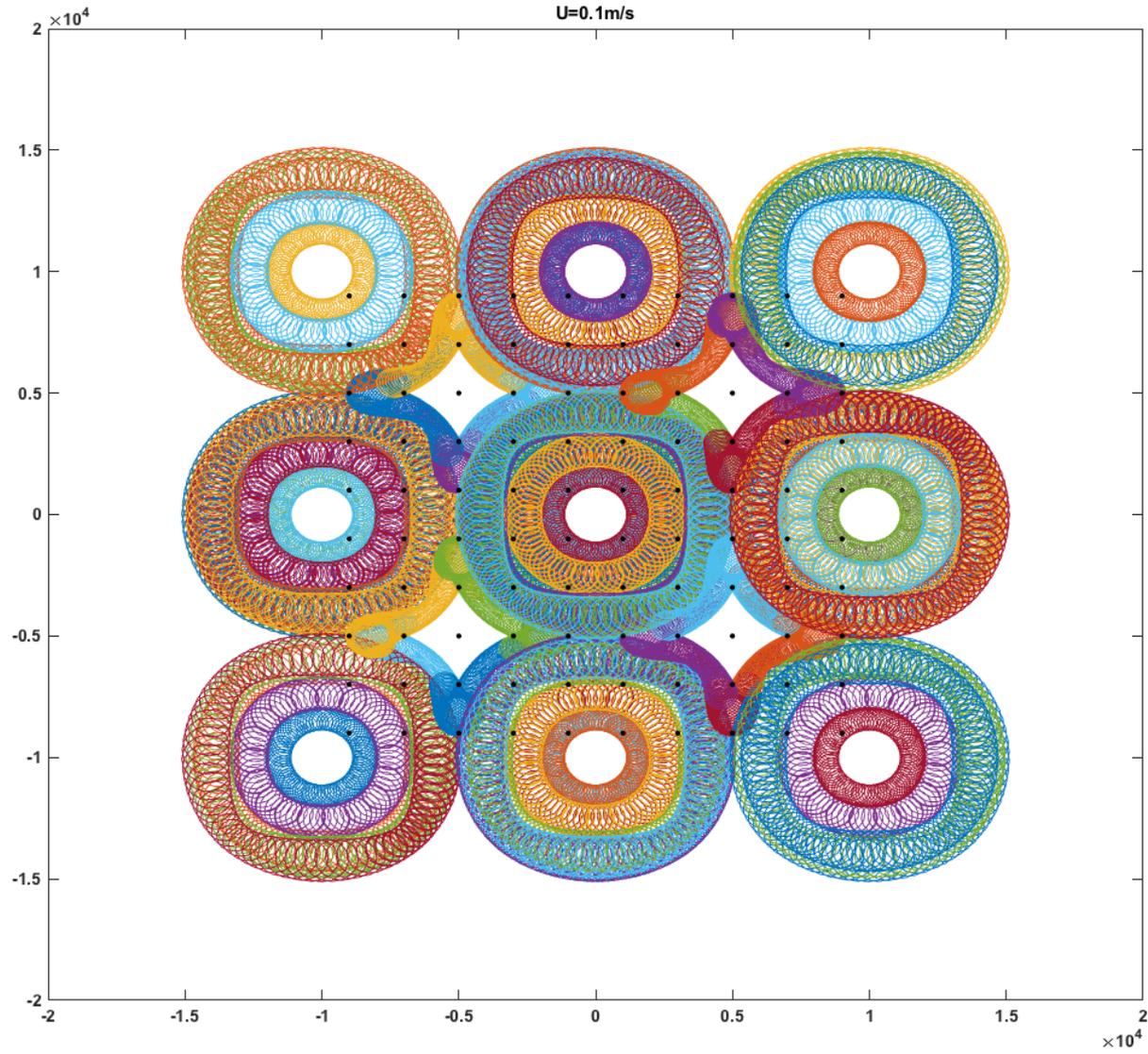
# Recap

- One flow field as seen by 2 observers
- Internal wave solution with 4 well-defined transport boundaries
- Boundaries organized by critical trajectories that oscillate at internal wave frequency
- All other trajectories live on ‘strange’ tori with highly variable return periods
- Consequence: aperiodic mixing within each transport region — distinct from traditional IW diapycnal mixing

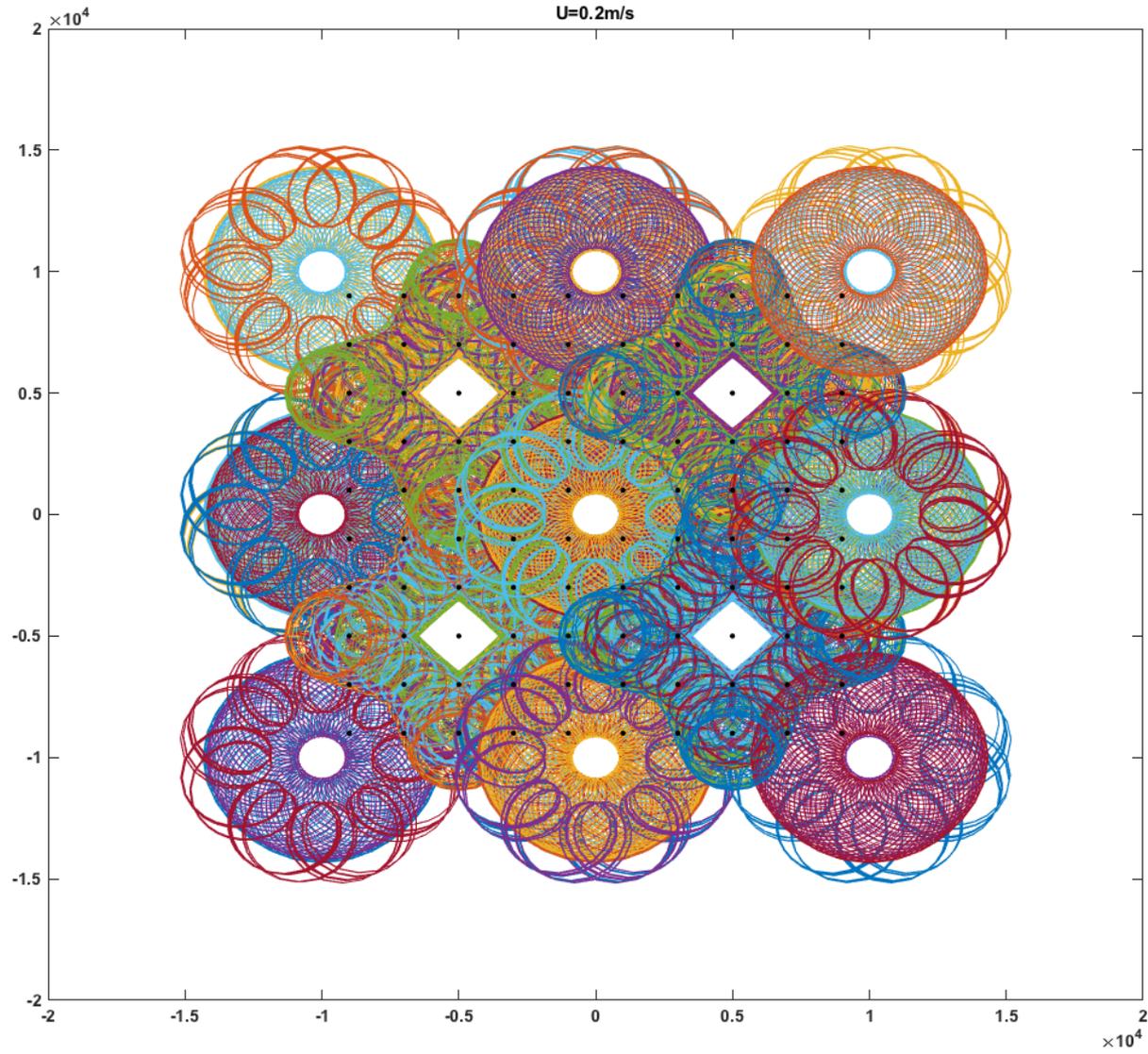
# Sidebar – Another Solution

- Waves occur at many different scales in oceanography
- Breaking waves are common in oceanography
- What happens when particle velocity exceeds phase velocity of disturbance?

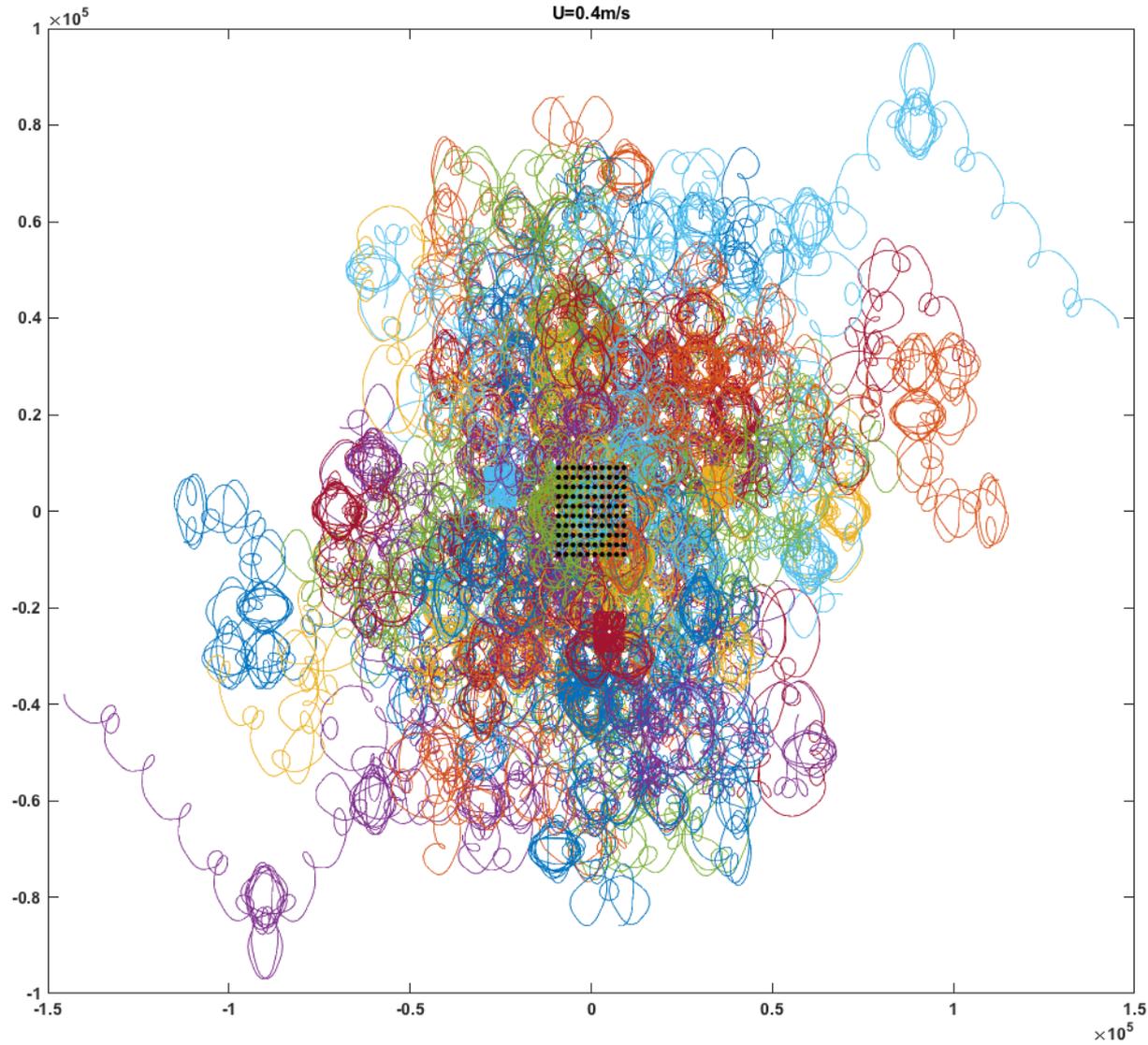
$$C_p = 35 \text{ cm/s}, U_W = 10 \text{ cm/s}$$



$$C_p = 35 \text{ cm/s}, U_W = 20 \text{ cm/s}$$



# Incoherent Mixing: $C_p = 35$ cm/s, $U_W = 40$ cm/s



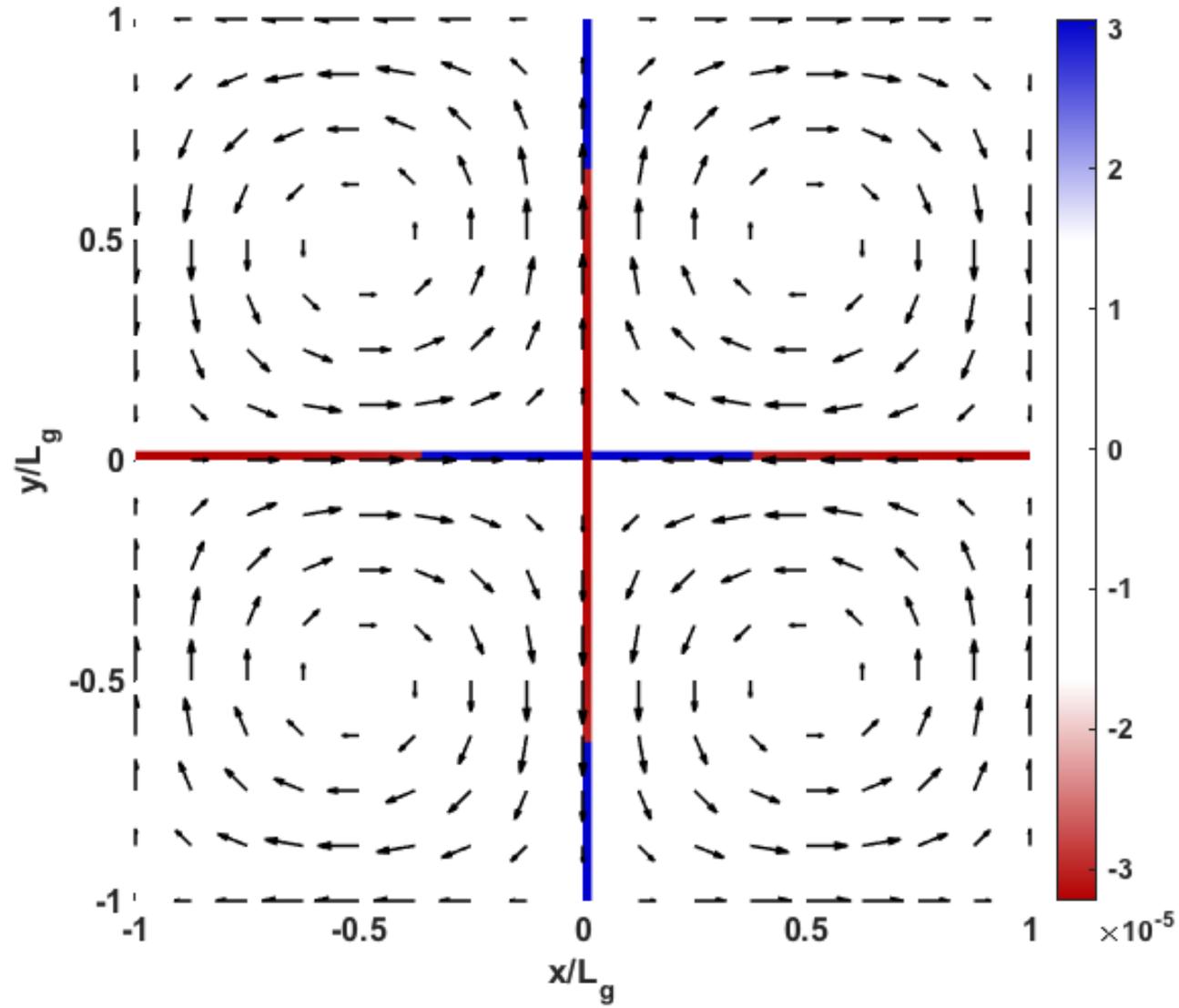
---

# **Can Internal Waves Disrupt Mesoscale Transport Pathways?\***

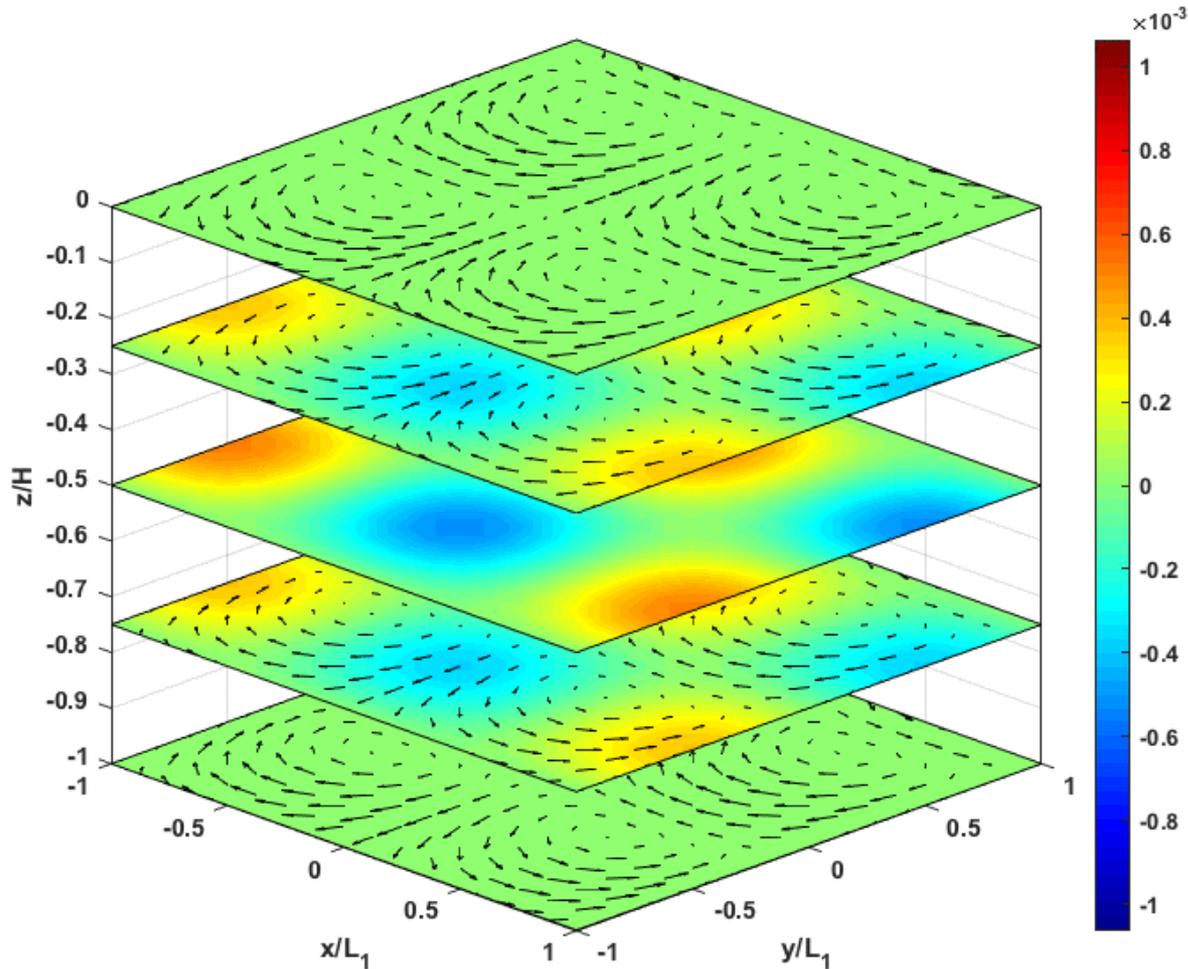
\*Chang et al., MS 162 Superior B 2:15 today:  
A Stratified 3D Model for Ocean Flows  
with Internal Waves

---

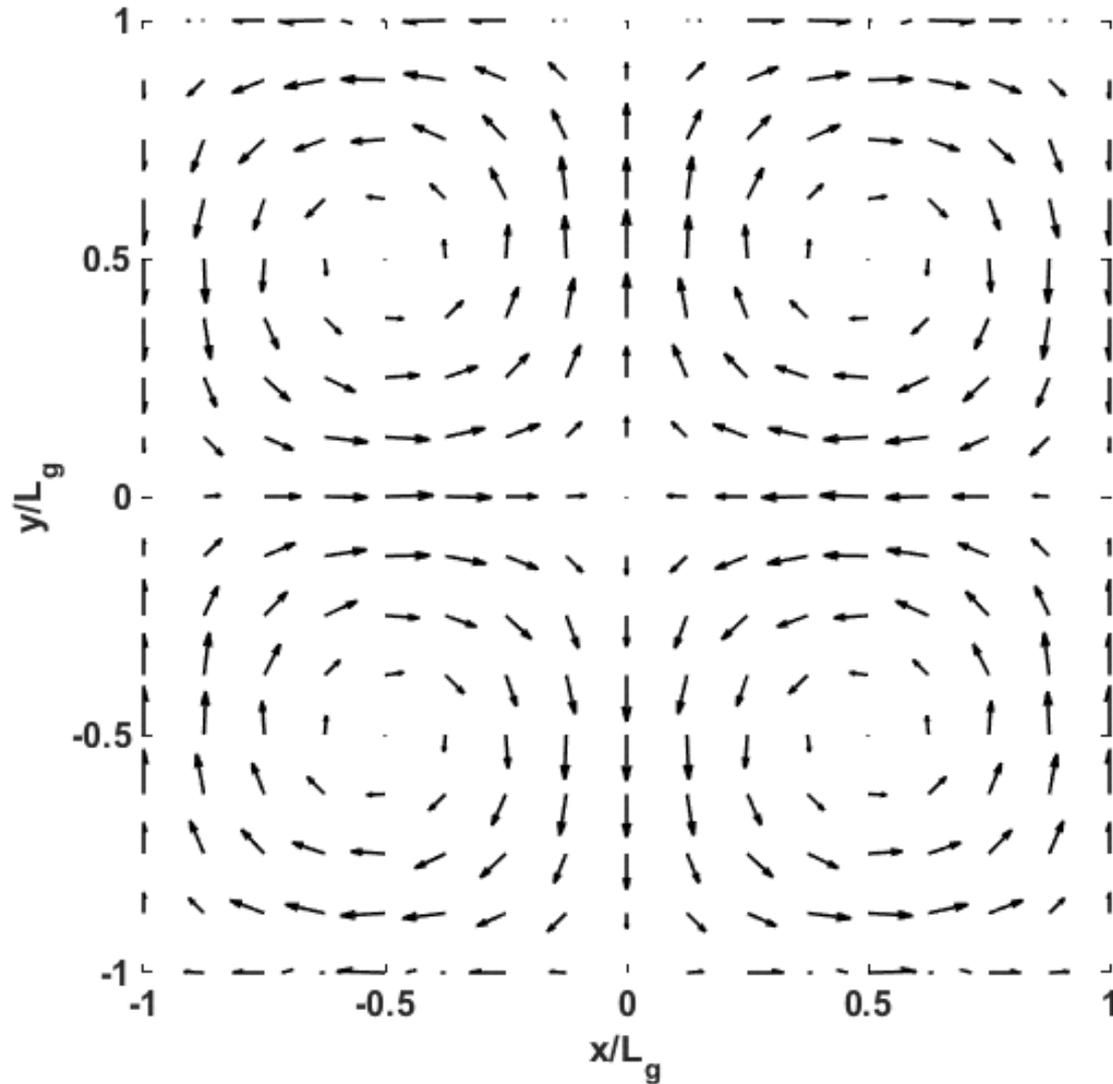
# Surface Velocity & FTLE for QP



# Traveling Internal Wave Velocity

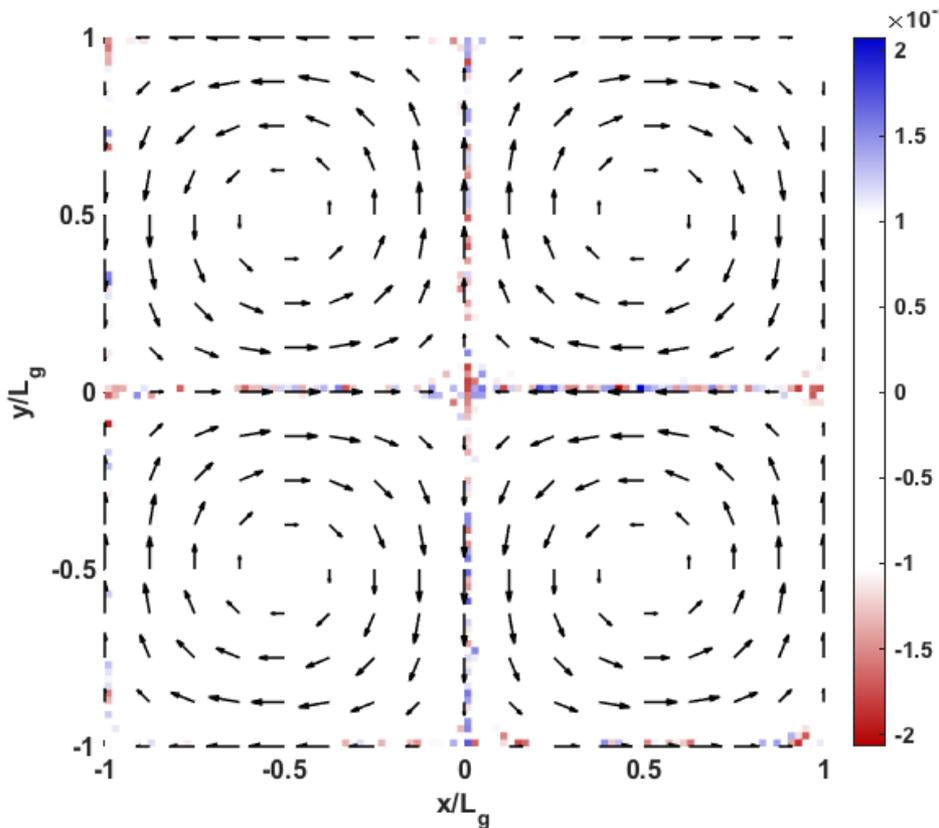


# QP + IW Surface Velocity

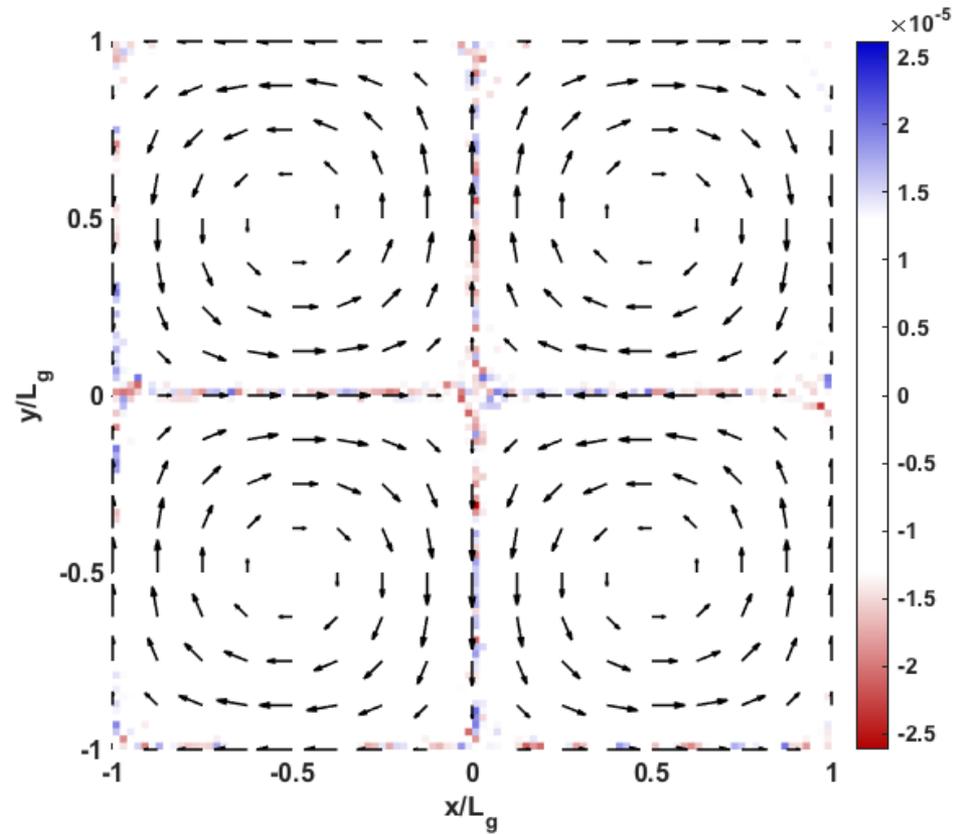


# FTLE for QP and Wave

## Surface

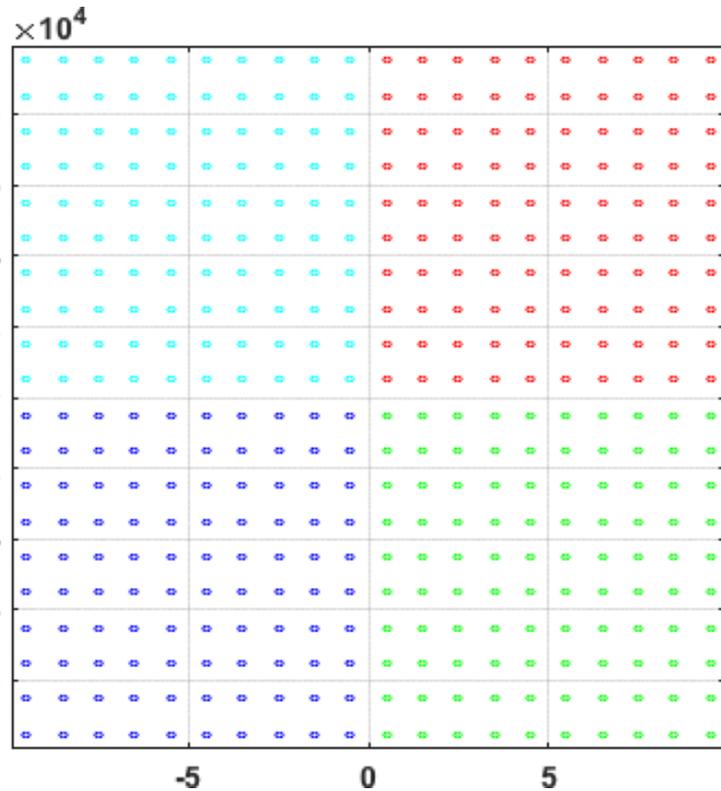


## Subsurface

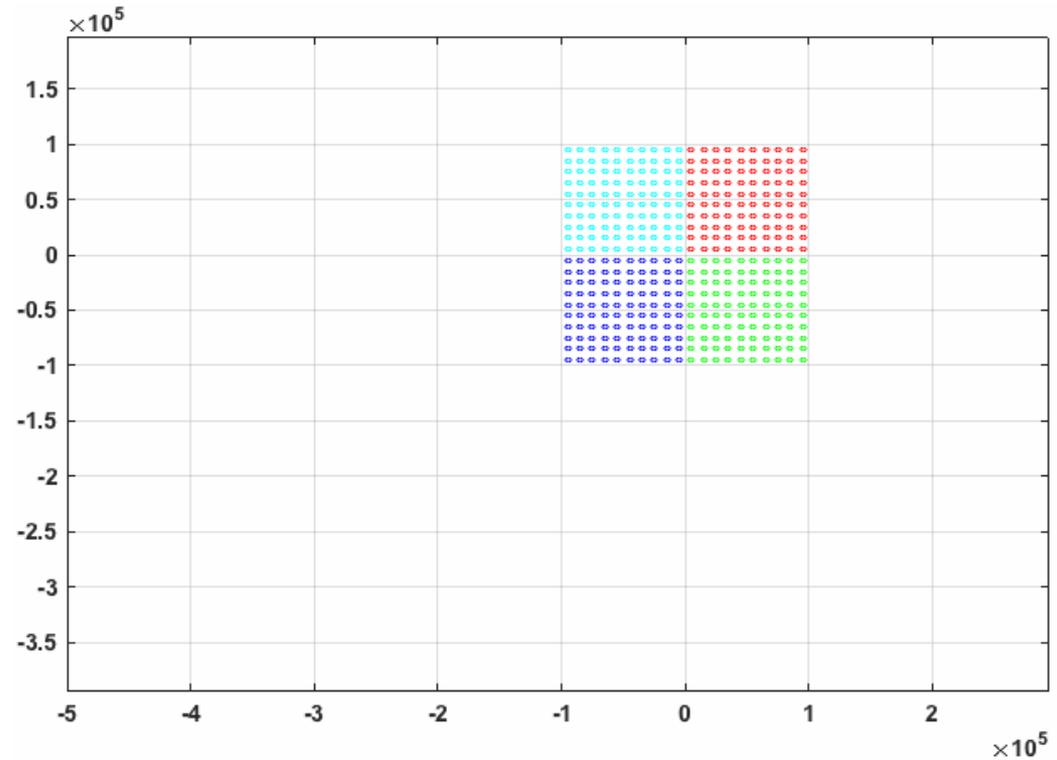


# Effect of IW Disturbance on QP

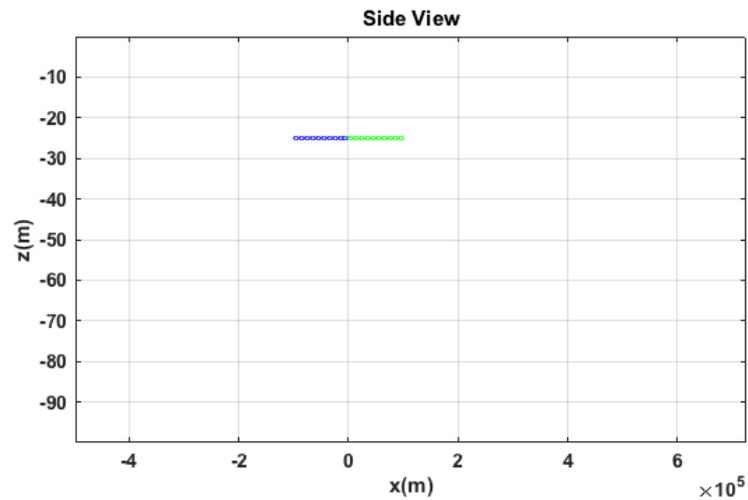
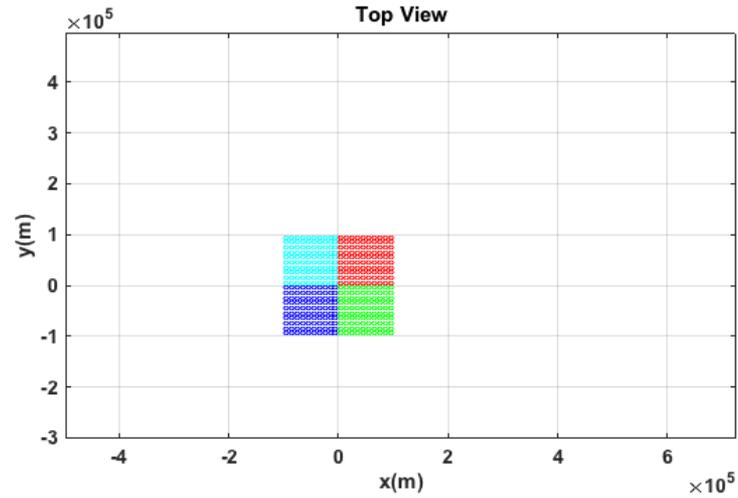
Just QP



QP + IW



# QP and IW – Mixing at H/4



# Summary

- Example of objective diagnostics in analyzing time dependent transport *surfaces*
- Internal waves may have well-defined ephemeral transport boundaries identified by 3D velocity
- Aperiodic mixing in internal waves distinct from diapycnal mixing
- Incoherent mixing when particle velocity exceeds phase velocity of traveling disturbance
- Internal waves cause leakage across mesoscale transport boundary surfaces

# Envoi

Hypothesis: Internal waves cause turnstile exchange across QP boundaries. Since this is a 3D flow field the turnstiles are 2D surfaces. It would be interesting to see what they look like.

Thanks for your time and patience

---

# Header

